

## The Tithonian Chitinoideidae and other microfossils from Owadów–Brzezinki quarry (central Poland)

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**Key words:** Tithonian, microfossils, Central Poland.

**Abstract.** Tithonian (= “Middle Volgian”) carbonate rocks are exposed in Owadów–Brzezinki quarry 19 km southeast of Tomaszów Mazowiecki, central Poland. In the upper part of the Sławno Limestone Member, chitinoideids have been identified in thin sections from three samples, only. Therefore, the documented part of the Chitinoideida Zone in Owadów–Brzezinki quarry is about 0.3 m thick. The identified chitinoideid taxa suggest that this assemblage represents the Upper Tithonian Boneti Subzone. The Chitinoideida Zone occurs at the top of Unit I and correlates with the uppermost interval of the Zarajskensis Horizon (Matyja, Wierzbowski, 2016). Other microfossils identified in the Chitinoideida Zone consist of *Saccocoma* sp. and benthic foraminifera of the genus *Planularia*.

Calcareous dinoflagellate cysts (*Cadosina semiradiata semiradiata* Wanner and *C. cf. semiradiata semiradiata* Wanner) occur above the chitinoideid assemblage, in the strata corresponding to the Gerassimovi Subzone of the Virgatus Zone (Matyja, Wierzbowski, 2016). Calcareous nannofossils are extremely rare in the thin sections studied; only one small specimen was seen, identified as *Rhombolithon minutum* (Rood *et al.*, 1971) Young *et Bown* 2014. In contrast, microbial filaments are frequently observed in the studied thin sections. Their shape and pattern are reminiscent of some recent Cyanobacteria of the order Nostocales, however the Tithonian microbial filaments are much thinner.

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### INTRODUCTION

Tithonian (= “Middle Volgian”) carbonate sediments are exposed in Owadów–Brzezinki quarry 19 km southeast of Tomaszów Mazowiecki, central Poland (Kin, Błażejowski, 2012; Kin *et al.*, 2013; Błażejowski *et al.*, 2014; Fig. 1). The section, about 32 m thick, exposed in this quarry (Fig. 2) belongs to the Pałuki Formation (topmost part) and Kcynia Formation (Matyja, Wierzbowski, 2016). The chalky limestones of the lower part of the Kcynia Formation (11.3 m) are distinguished as the Sławno Limestone Member (Matyja, Wierzbowski, 2016). The bulk of these limestones belongs to the Scythicus Zone, Zarajskensis Subzone and Horizon (Matyja, Wierzbowski, 2016). The identified chitino-

ideid taxa suggest that this assemblage may represent the Boneti Subzone. Other microfossils identified in the Chitinoideida Zone belong to *Saccocoma* sp. (ossicles) and benthic foraminifera.

### MATERIAL FOR STUDY

Dr. Błażej Błażejowski collected the samples used for the present study (Fig. 2). I have examined under the light microscope over 200 thin sections made from 117 samples but of these only 18 samples contained microfossils identified and reported herein (Fig. 2). Many thin sections contain ostracods only, or are barren. The study and microphoto-

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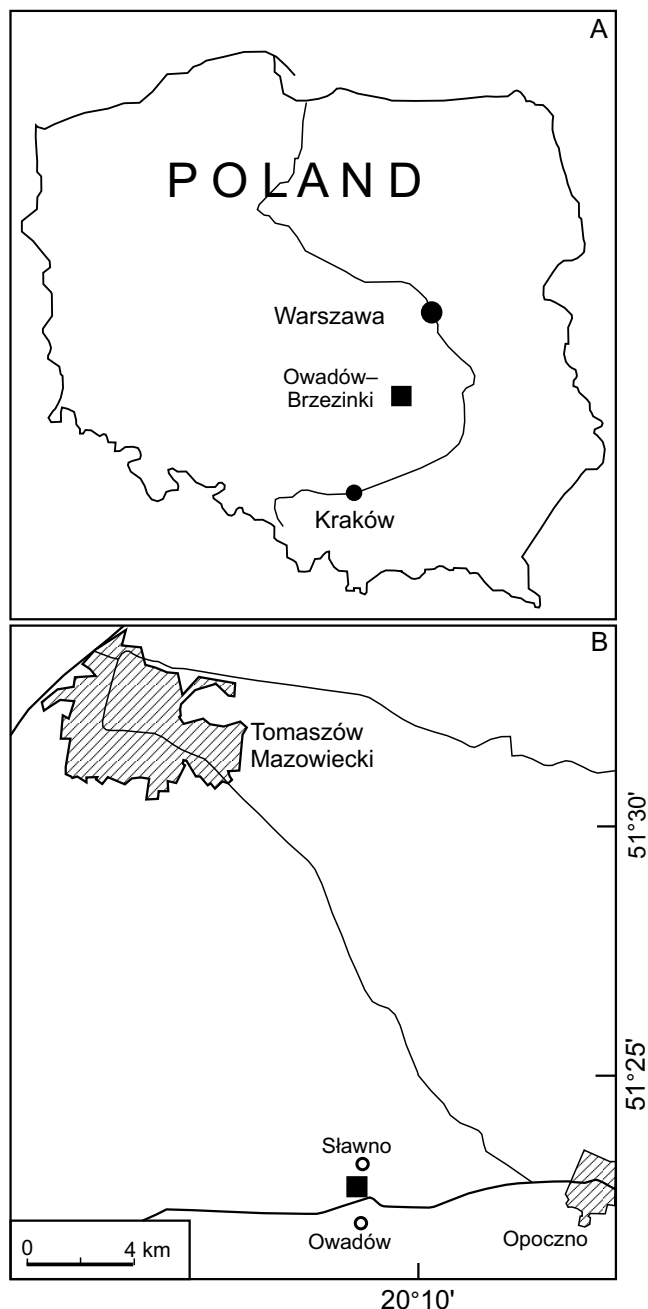


Fig. 1. A. Location of Owadów-Brzezinki quarry in Poland (after Kin *et al.*, 2012, fig. 1, modified). B. Location map of Owadów-Brzezinki quarry south of Sławno (partly after Salamon *et al.*, 2006, fig. 1B)

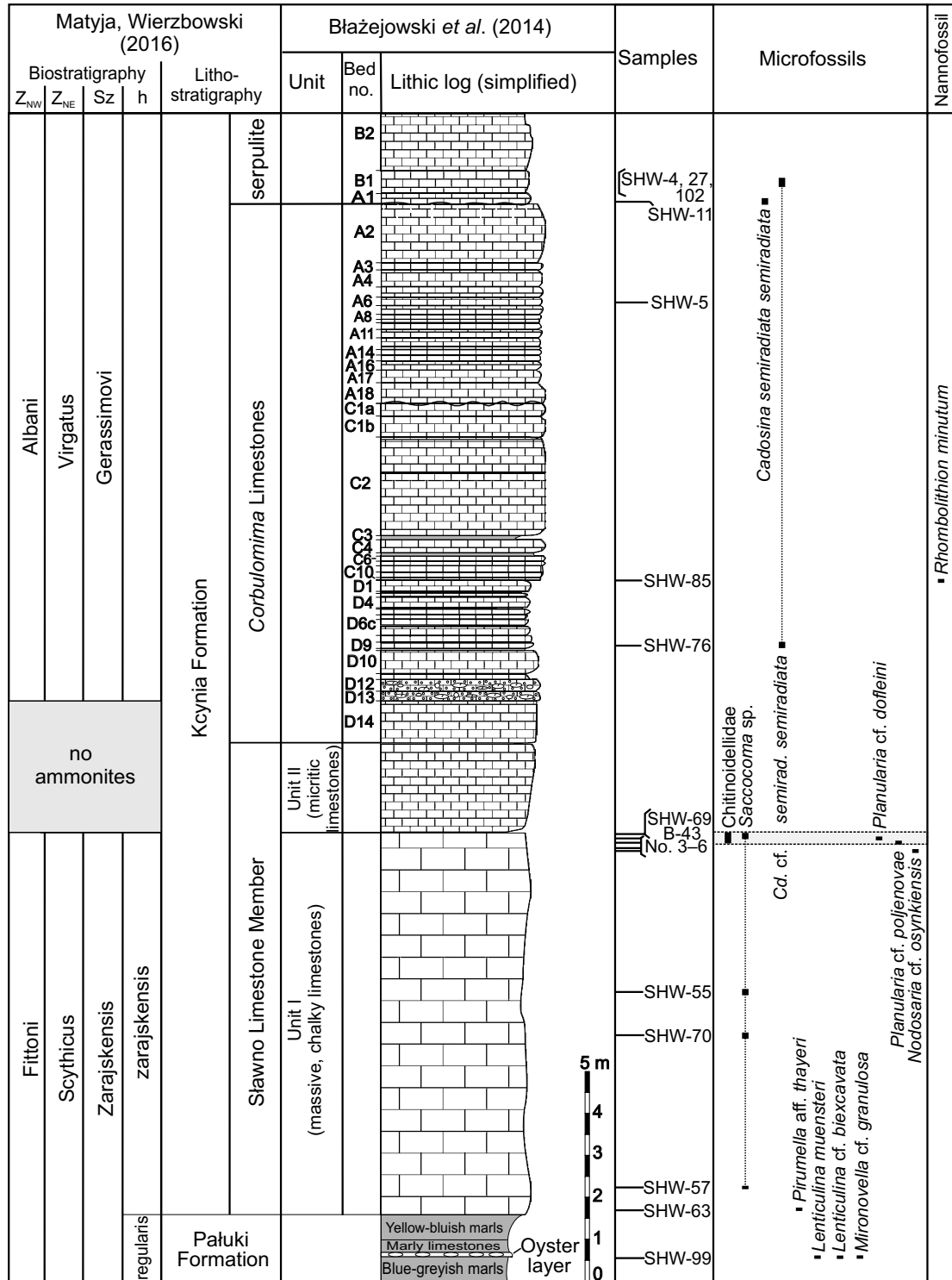
graphs of the microfossils from thin sections were made under a NIKON Polarizing Microscope (ECLIPSE LV100POL) in the Institute of Geological Sciences (Research Centre in Warsaw) of the Polish Academy of Sciences.

## RESULTS

Only a few benthic foraminifera were found in the uppermost part of the Pałuki Formation (Fig. 2). The following taxa occur in sample SHW-99: *Mironovella* cf. *granulosa* (Bielecka et Pożaryski), *Lenticulina muensteri* (Roemer) and *Lenticulina* cf. *biexcavata* (Mjatluk) (Fig. 5A–C). Chitinoideidellids have been identified in the thin sections corresponding to three samples (SHW-69, B-43 and no. 4) collected from the upper part of the Sławno Limestone Member (Kcynia Formation). Therefore, the documented part of the Chitinoideidella Zone in Owadów-Brzezinki quarry is about 0.3 m thick, only (Fig. 2). This biozone occurs at the top of the Unit I (massive chalky limestones of Błazejowski *et al.*, 2014), and correlates with the upper part of the Zarajskensis Subzone (Matyja *et al.*, 2016) and Horizon documented in the lower part of the section (Matyja, Wierzbowski, 2016). The chitinoideidellid taxa found comprise: *Borziella* cf. *slovenica* (Borza, 1969) (Fig. 3B, C), *Chitinoideidella* cf. *elongata* Pop, 1997 (Fig. 4A), *Chitinoideidella* aff. *boneti* Doben, 1963 (Fig. 4B), *Daciella* cf. *banatica* Pop, 1998 (Fig. 3D), *Daciella* sp. (Fig. 3A) and *Dobeniella* sp. (Fig. 3E, F). A mould of a chitinoideidellid similar to *Chitinoideidella boneti* Doben, 1963 (Fig. 4C) was found in thin section no. 4. The identified chitinoideidellid taxa suggest that this assemblage represents the Boneti Subzone. Other microfossils identified in the Chitinoideidella Zone belong to *Saccocoma* sp. (scarce ossicles) and benthic foraminifera: *Planularia* cf. *poljenovae* (Kasanzev) and *P.* cf. *dofleini* Kusnetzova. Sporadic ossicles of *Saccocoma* sp. occur in the Sławno Limestone Member (Fig. 4D, E).

In the studied samples calcareous dinoflagellate cysts (*Cadosina semiradiata semiradiata* Wanner and *C.* cf. *semiradiata semiradiata* Wanner) occur above the chitinoideidellid assemblage (Fig. 2), in the strata corresponding to the Gerassimovi Subzone of the Virgatus Zone (Matyja *et al.*, 2016; Matyja, Wierzbowski, 2016). This upper part of the section correlates with the Albani Zone of the mid Tithonian in England and Northern France (Matyja *et al.*, 2016; Matyja, Wierzbowski, 2016). A specimen of *Pirumella* aff. *thayeri* (Bolli, 1974) Lentini et Williams (1993) was recorded at the base of the Sławno Limestone Member, well below the chitinoideidellid assemblage (Fig. 2). In the Brodno section (Pieniń Klippen Belt, Slovakia), *C. tenuis* (included in the taxon *Pirumella thayeri* by Ivanova, Keupp, 1999) was recorded in the Dobeni and Boneti subzones of the Chitinoideidella Zone, up to the Remanei Subzone of the Crassicolliaria Zone (Michalik *et al.*, 2009, fig. 8).

Rogov (2014, fig. 1) correlated the Zarajskensis Subzone of the Scythicus Zone and the Gerassimovi Subzone of the



D12–D13: Fossiliferous beds with horseshoe crab fauna

----- Chitinoideiella Zone

Fig. 2. Stratigraphical section exposed in Owadów–Brzezinki quarry. Ammonite biostratigraphy after Matyja, Wierzbowski (2016)

Z<sub>NW</sub> – zones of the NW part of the Subboreal Province, Z<sub>NE</sub> – zones of the NE part of the Subboreal Province, Sz – Subzones, h – horizons; samples location after B. Błażejowski (written information)

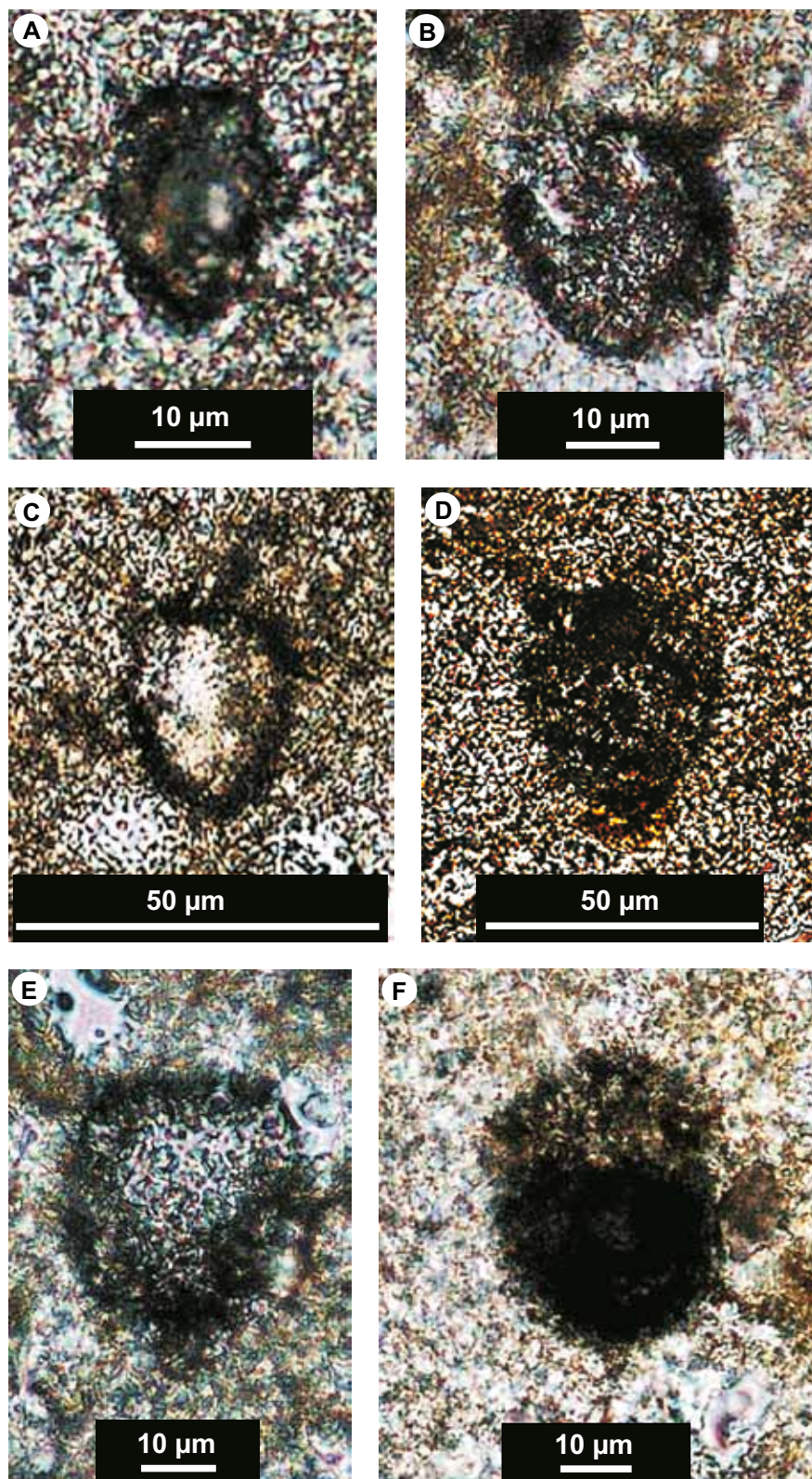


Fig. 3. Chitinoidellids identified in the sample/thin section SHW-69 (for location see Fig. 2)

A. *Daciella* sp. B., C. *Borziella* cf. *slovenica* (Borza, 1969). D. *Daciella* cf. *banatica* Pop, 1998. E., F. *Dobeniella* sp.

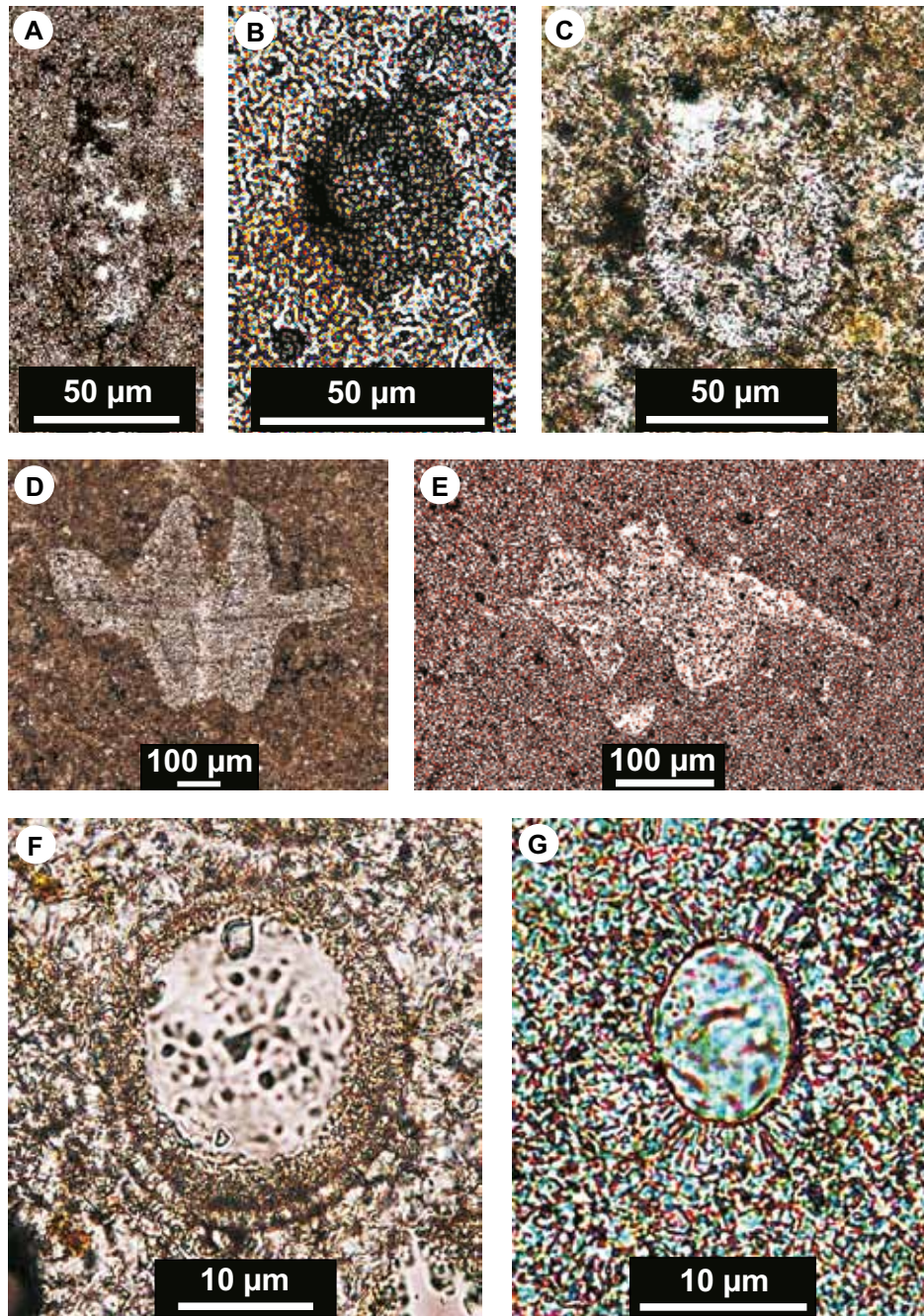


Fig. 4. Chitinoideiellids, *Saccocoma*, *Cadosina* and *Pirumella*

**A.** *Chitinoidella* cf. *elongata* Pop, 1997, sample SHW-69, thin section B-43. **B.** *Chitinoidella* aff. *boneti* Doben, sample SHW-69, thin section B-43. **C.** Mould similar to *Chitinoidella boneti* Doben, 1963 (thin section no. 4). **D.** Ossicle of *Saccocoma* sp., sample SHW-70. **E.** Ossicle of *Saccocoma* sp., sample/thin section SHW-57. **F.** *Cadosina semiradiata semiradiata* Wanner, 1940, sample/thin section SHW-11. **G.** *Pirumella* aff. *thayeri* (Bolli, 1974) Lentin et Williams, 1993, sample SHW-63

Virgatus Zone of the Russian Platform with the upper part of the Lower Tithonian Ponti Zone (in Spain). If this correlation is correct, the chitinoideiellid assemblage from Owadów–Brzezinki quarry should correspond to the upper part of the

Ponti Zone in Spain. However, in the Puerto Escaño section (Spain), Pruner *et al.* (2010) have found the Dobeni/Boneti subzonal boundary in the Upper Tithonian Simplisphinctes Zone. Lakova, Petrova (2013) correlate the Dobeni/Boneti

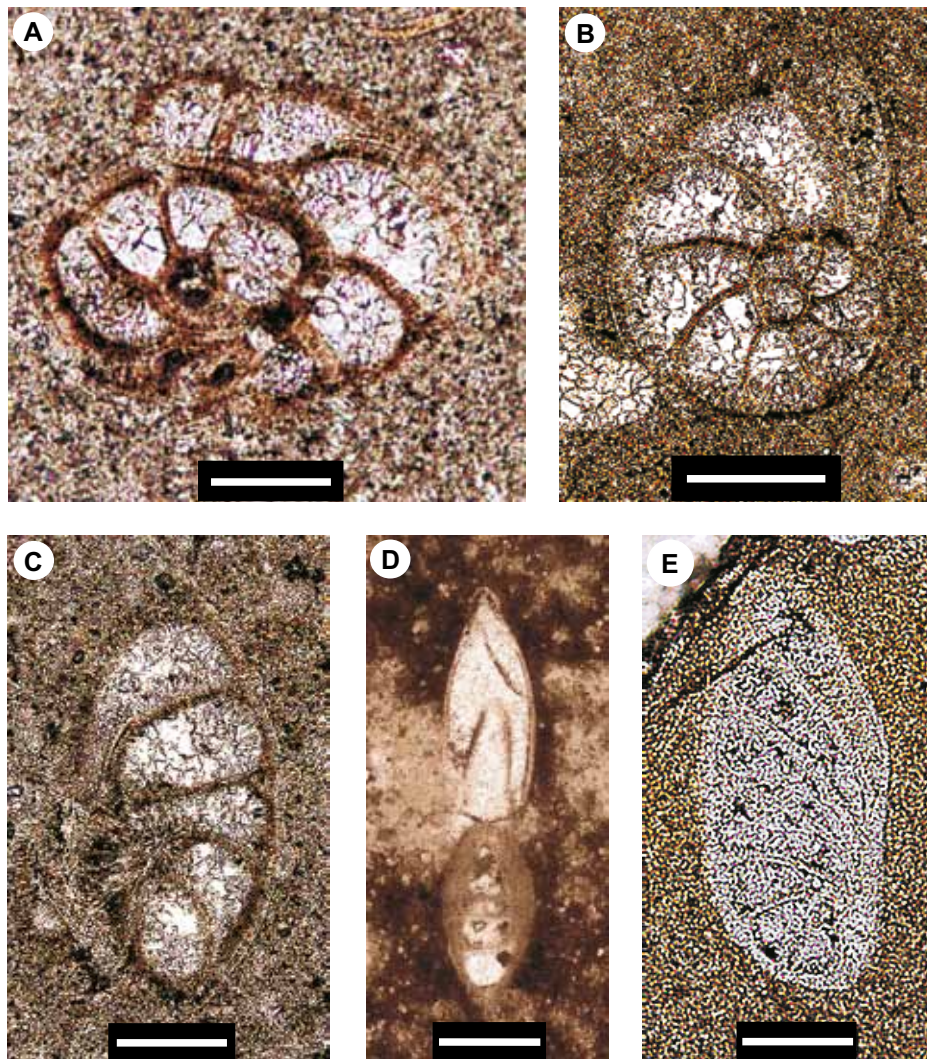


Fig. 5. Benthic foraminifera from the Pałuki Formation (A–C) and the Sławno Limestone Member (D–E)

A. *Mironovella* cf. *granulosa* (Bielecka et Pożaryski, 1954), sample/thin section SHW-99. B. *Lenticulina muensteri* (Roemer), sample/thin section SHW-99. C. *Lenticulina* cf. *biexcavata* (Mjatluk, 1939), sample/thin section SHW-99. D. *Planularia* cf. *dofleini* (Kasanzev, 1936), additional thin section no. 3. E. *Planularia* cf. *poljenovae* Kusnetzova, 1960 (additional thin section no. 4). Scale bar – 100  $\mu$ m

subzonal boundary with the base of the Upper Tithonian. Therefore, the uppermost part of the Zarajskensis Subzone and the Gerassimovi Subzone in Owadów–Brzezinki quarry may belong to the Upper Tithonian.

## CONCLUSIONS

In Owadów–Brzezinki quarry (Central Poland), the following benthic foraminifera occur in the uppermost part of the Pałuki Formation: *Lenticulina muensteri*, *L.* cf. *biexcavata* and *Mironovella* cf. *granulosa*.

Chitinoideidae have been identified in the upper part of the Sławno Limestone Member (Kcynia Formation). The taxa found include: *Borziella* cf. *slovenica* (Borza), *Chitinoideella* cf. *elongata* Pop, *Chitinoideella* aff. *boneti* Doben, *Daciella* cf. *banatica* Pop, *Daciella* sp. and *Dobeniella* sp. A mould similar to *Chitinoideella boneti* Doben was also found.

The documented part of the Chitinoideella Zone only is about 0.3 m thick. The chitinoideid taxa indicate the lower part of the Boneti Subzone. Other microfossils identified in the Chitinoideella Zone belong to *Saccocoma* sp. (ossicles) and benthic foraminifera: *Planularia* cf. *poljenovae* (Kasanzev) and *P.* cf. *dofleini* Kusnetzova.

The Boneti Subzone of the Chitinoideidella Zone found in the quarry correlates with the upper part of the Zarajskensis Subzone and Horizon recognized in the lower part of the section (Matyja, Wierzbowski, 2016). Therefore, the upper part of the Zarajskensis Subzone of the Scythicus Zone and the Gerassimovi Subzone of the Virgatus Zone may belong to the Upper Tithonian.

The calcareous dinoflagellate cysts *Cadosina semiradiata semiradiata* Wanner and *C. cf. semiradiata semiradiata* Wanner occur above the chitinoideidellid assemblage, in the strata corresponding to the Gerassimovi Subzone (Matyja *et al.*, 2016; Matyja, Wierzbowski, 2016). One specimen of *Pirumella* aff. *thayeri* (Bolli, 1974) Lentini et Williams (1993) (or *Colomisphaera* aff. *tenuis* [Nagy, 1966]) was recorded at the base of the Sławno Limestone Member.

## SYSTEMATIC NOTES

### Chitinoideidellids

*Borziella* cf. *slovenica* (Borza, 1969)

Fig. 3B, C

**Description:** Two small specimens with an oval (B) to almost subtriangular lorica (C). The specimens are 27 µm and 32 µm in length, 24.8 and 22.5 µm in width, respectively. The collar is rather short, funnel-like.

**Remarks:** The specimen shown in Figure 3B is similar to the holotype of *B. slovenica* (Borza, 1969, pl. LXVI, fig. 9). The other specimen (Fig. 3C) is similar to *B. slovenica* figured by Borza (1969, pl. LXVI, fig. 12). According to Reháková (2002), *B. slovenica* (Borza, 1969) occurs in the Dobeni Subzone of the Tithonian Chitinoideidella Zone.

*Chitinoideidella* cf. *elongata* Pop, 1997

Fig. 4A

**Description:** A poorly preserved specimen with an outwardly deflected collar; the lorica is about 84 µm long and 41 µm wide. The caudal appendage is not visible.

**Remarks:** The shape and dimensions of the specimen are typical of *Ch. elongata* Pop (cf. Pop, 1997 and Reháková, 2002), but the details of its structure are obliterated.

*Chitinoideidella* aff. *boneti* Doben, 1963

Fig. 4B

**Description:** Bell-shaped lorica, 39 µm long and 35.7 µm wide, with a preoral constriction. The length/width ratio is about 1.1. The aboral end of the lorica is sharp, with a short caudal appendage.

**Remarks:** The figured specimen is similar to *Chitinoideidella boneti* Doben, but is smaller in comparison with the

dimensions reported for this species (Borza, 1969; Reháková, 2002).

*Daciella* cf. *banatica* Pop, 1998

Fig. 3D

**Description:** The figured specimen (46.1 µm long and 37.3 µm wide) has a conical lorica (38.8 µm in length) sharply terminated at the aboral pole, but the caudal appendage is not visible. The outwardly deflected collar seems to be as wide as the maximum width of the lorica.

**Remarks:** The rather poor preservation of the specimen does not allow its full identification to species level. It is similar to the holotype (Pop, 1998, fig. 1.2), excepting the caudal appendage, which is not observed in this section.

*Daciella* sp.

Fig. 3A

**Description:** A small specimen, 21.5 µm long and 14.3 µm wide, with a conical and slightly elongate lorica. The aboral pole is pointed but the caudal appendage is (almost) invisible. The oral end of the lorica shows a characteristic constriction 1.3 µm wide. The rather large oral opening bears a short collar preceded by a distinct swelling.

**Remarks:** The specimen is similar to *Daciella banatica* n. sp. (Pop, 1998, fig. 2: 8–9) and also to *Chitinoideidella* sp. illustrated by Borza (1969, pl. LXIX, fig. 6) and included by Pop (1998) in the synonymy of *D. banatica* n. sp. However, the specimen figured herein is smaller than the holotype and paratype of *D. banatica* Pop, and the caudal appendage is hardly visible in Figure 3A. According to Pop (1998), the species of *Daciella* occur in the Chitinoideidella and Praetintinnopsella zones (Tithonian).

*Dobeniella* sp.

Fig. 3E, F

**Description:** The specimen with a bell-shaped lorica and short aboral horn (Fig. 3E) is 36 µm long and 34 µm wide. The collar is hardly visible, but is similar to that of *D. colomi* (Borza, 1966). The lorica is filled with fossilized microbial filaments, also overgrowing the aboral horn. The specimen shown in Fig. 3F is 40.4 µm long and 28 µm wide, with an ovoid lorica about 24 µm long. A constriction separates the lorica from the collar, which is poorly visible because of microbial filaments filling the specimen. The caudal appendage is sharply pointed, about 7 µm long.

**Remark:** Both chitinoideidellids are poorly preserved, but their shape is similar to *Dobeniella* Pop, 1997. The specimen shown in Fig. 3F cannot be identified to species rank because the structure of the collar is not fully discernible.

## Saccocomid ossicles

*Saccocoma* sp.  
Fig. 4D, E

**Remarks:** The figured fragments of *Saccocoma* sp. are small skeletal sections observed in two different thin sections (from sample SHW-70 and SHW-57). The first one (Fig. 4D) shows a section similar to the second primibrachials of *S. aff. vernioryi* Manni & Nicosia (Brodacki, 2006, fig. 4E<sub>4</sub>). The saccocomid fragment shown in Fig. 4E is similar to the skeletal element BI (Benzaggagh *et al.*, 2015, fig. 4) described therein as “irregular heads with one or two foofs and two lateral appendices (irg. Hd)”. In the stratigraphical scheme for the western Tethyan realm (Benzaggagh *et al.*, 2015, fig. 17) the skeletal sections, labelled BE in their figure 17, are abundant in saccocomid biozone 4, which corresponds to the ammonite zones Darwini–Semiforme (Lower Tithonian).

## Calcareous dinoflagellates

*Cadosina cf. semiradiata semiradiata* Wanner, 1940  
Fig. 4F

**Remarks:** The figured specimen found in the sample SHW-4 (54 µm in diameter) is similar to *Cadosina semiradiata semiradiata* Wanner, 1940 (*cf.* Ivanova, Keupp, 1999, pl. I, figs 6–7; Reháková, 2000a, pl. II, fig. 6). The wall of the test (about 11 µm) consists of a dark inner layer (7 µm) and a light grey outer one (4 µm). The calcareous dinoflagellate *Cadosina cf. semiradiata semiradiata* Wanner was also found in other samples (SHW: 11, 27, 76 and 102). Ivanova, Keupp (1999) included *Cadosina semiradiata semiradiata* in the species *Pirumella multiestrata* (Pflaumann et Krashe-ninnikov, 1978) Lentin et Williams, 1993 from the subfamily Obliquipithonelloideae Keupp, 1987. Reháková (2000a, b) placed the Semiradiata Zone in the Middle Tithonian.

*Pirumella aff. thayeri* (Bolli, 1974) Lentin et Williams 1993  
Fig. 4G

**Remarks:** The small specimen identified as *Pirumella aff. thayeri* is ovate in shape. The maximal diameter of the test is 18 µm only, whereas the wall thickness attains 3.7 µm. According to Bolli (1980), the tests of the genus *Pirumella* are circular to ovate in axial section. Although very small, the figured specimen is partly similar to *Colomisphaera tenuis* (Nagy, 1966) reported by Ivanova, Keupp (1999) from the base of the Chitinoidella Zone (Middle Tithonian) to the Lower Berriasian C. alpina Subzone. Řehánek (1992, fig. 2) placed the Tenuis Zone at the Lower/Upper Tithonian boundary. According to Reháková (2000a), *C. tenuis* ranges from the Middle Tithonian to the Upper Berriasian.

## Benthic foraminifera

*Mironovella cf. granulosa* (Bielecka et Pożaryski, 1954)  
Fig. 5A

**Remarks:** The figured specimen, found in the sample/thin section SHW-99, is similar to the holotype illustrated by Bielecka, Pożaryski (1954, pl. XII, fig. 61b) as *Epistomina stelicostata* var. *granulosa*. *Mironovella granulosa* was reported to occur from the Late Oxfordian to Early Tithonian in the northern epicontinental areas of the Tethys (Görög, Wernli, 2013).

*Lenticulina muensteri* (Roemer, 1839)  
Fig. 5B

1839. *Robulina Münsteri* Roemer; Roemer, p. 48, pl. 20, fig. 29 (*fide* Cat. of Foram. Ellis & Messina, 1940) – *vide* Jendryka-Fuglewicz (1975) (see Jendryka-Fuglewicz, 1975 for synonymy up to 1975).

1975. *Lenticulina muensteri* (Roemer); Jendryka-Fuglewicz, p. 149, pl. VIII–X, pl. XI, figs 1–6, pl. XIX, pl. XX, figs 1, 2.

1983. *Lenticulina muensteri* (Roemer, 1839); Peryt, p. 440, pl. 24: 7, 8.

1986. *Lenticulina muensteri* (Roemer); Bartenstein & Bolli, pl. 4, figs 25, 26.

1995. *Lenticulina muensteri* (Roemer); Holbourn & Kaminski, p. 214, pl. 6, figs 11a, b.

**Description:** The specimen has an involute test with 8 chambers in the last whorl; the sutures are markedly curved backward. The maximum diameter of the test attains 282 µm.

**Remarks:** The figured section, found in the sample/thin section SHW-99, is similar to the morphotype *wisniowskii* of *L. muensteri* occurring in the Middle and Upper Jurassic (Jendryka-Fuglewicz, 1975).

*Lenticulina cf. biexcavata* (Mjatluk, 1939)  
Fig. 5C

**Remarks:** The specimen is an equatorial section 394 µm in length, partly overgrown with microbial filaments; therefore, only 5 chambers are visible in the outer whorl. The sutures are usually arcuate and depressed. The specimen was found in the sample/thin section SHW-99 and is similar to tests illustrated by Jendryka-Fuglewicz (1975, pl. III, figs 5–8) being, however, more slender in the equatorial section. According to Jendryka-Fuglewicz (1975), *L. biexcavata* is known from the Middle Jurassic of Poland, Lower Kimmeridgian of France and “Lower Volgian” in Russia.

*Planularia cf. dofleini* (Kasanzev, 1936)  
Fig. 5D

**Remarks:** The figured peripheral section is broadly similar to the test of *Planularia dofleini* (Kasanzev) illustrated



by Bielecka (1975, pl. IX, fig. 1b). However, the specimen shown in Fig. 5D is much smaller (410  $\mu\text{m}$  in length). In Poland, *P. dofleini* (Kasanzev) was reported from the Zaraiskites scythicus and *Z. zaraiskensis* zones (“Middle Portlandian” – Bielecka, 1975, table III).

*Planularia* cf. *poljenovae* Kusnetzova, 1960

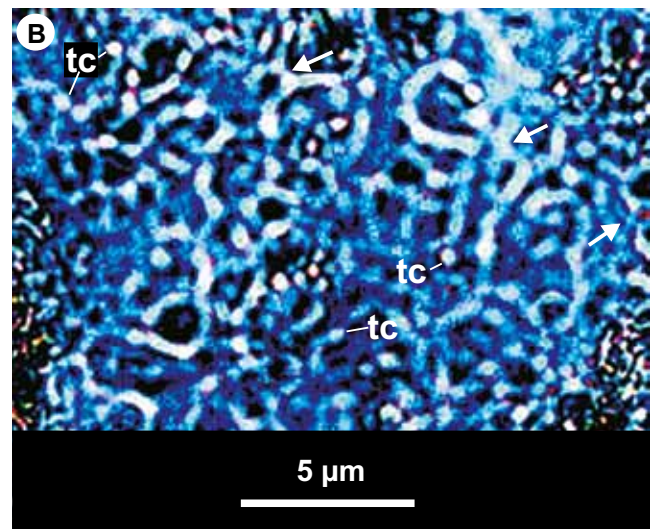
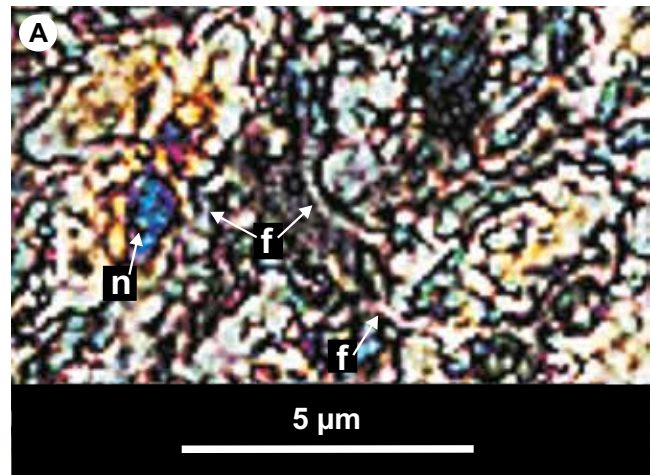
Fig. 5E

**Remarks:** The specimen identified as *Planularia* cf. *poljenovae* Kusnetzova, 376  $\mu\text{m}$  in length, is broadly similar to *P. poljenovae* illustrated by Bielecka (1975, pl. IX, figs 4–7), although its upper part seems to be less elongated than the tests shown in that publication. According to Bielecka (1975), *P. poljenovae* is known from the lower and middle “Portlandian” in Poland and the lower and middle “Volgian” on the Russian Platform.

## NANNOFOSSIL AND MICROBIAL FILAMENTS

Calcareous nannofossils are extremely rare in the studied thin sections; only one small specimen was identified (Fig. 6A, n) as *Rhombolithion minutum* (Rood *et al.*, 1971) Young et Bown 2014 (*cf.* Young *et al.*, 2015). Originally this taxon was described as *Diadorhombus minutus* Rood *et al.*, 1971 (see also Rood, Barnard, 1972) from the Oxfordian, however its stratigraphical range seems to be wider because De Kaenel, Bergen (1996) reported *Diadorhombus scutulatus* (Medd, 1971), Medd, 1979 from the Lower Tithonian at Hole 901A (Atlantic Ocean). According to Young *et al.* (2015), the taxa *R. minutum* and *R. scutulatus* (Medd, 1971) Young et Bown 2014 “are evidently synonyms it is not clear yet which has priority”.

The above described nannofossil is surrounded by microbial filaments 0.15–0.80  $\mu\text{m}$  wide and up to 4  $\mu\text{m}$  long (Fig. 6A, f). Although thinner, the filaments (usually branched) are similar to some recent Cyanobacteria from the order Nostocales (Komárek *et al.*, 2014). Microbial filaments are frequently observed in the studied thin sections. The filaments 0.2–0.6  $\mu\text{m}$  wide (Fig. 6B) probably infilling the space occupied by a microfossil, which no longer exists. Their shape and pattern are reminiscent of some recent Cyanobacteria from the family Hapalosiphonaceae, order Nostocales (Komárek *et al.*, 2014). The filaments reveal true branching; some terminal cells are similar to heterocysts (Fig. 6B, tc). However, the figured fossil filaments are much thinner than the recent Cyanobacteria. For example, the genus *Mastigocoleus* Lagerheim has trichomes 4–6  $\mu\text{m}$  wide (Wuitner, 1921, fig. 9; Carreiro-Silva *et al.*, 2012, fig. 2f). Some other genera of the family Hapalosiphonaceae, such as *Fischerella* Gomont, have trichomes about 3–6  $\mu\text{m}$  wide (*cf.* Rippka *et al.*, 1979, figs 73–78). Therefore, the figured



**Fig. 6.** A. Nannofossil (n) identified as *Rhombolithion minutum* (Rood *et al.*, 1971) Young et Bown 2014 and microbial filaments (f); *Corbulomima* Limestone Member, sample/thin section SHW-85; B. Microbial filaments (tc – terminal cells slightly similar to cyanobacterial heterocysts, but much smaller – 0.4–0.6  $\mu\text{m}$  wide; arrows indicate branching of filaments); sample/thin section SHW-5, *Corbulomima* Limestone Member, bed no. A6

microbial filaments (Fig. 6B) cannot be considered as the Tithonian equivalents of the above mentioned recent cyanobacterial taxa.

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