

The Bajocian to Kimmeridgian (Middle to Upper Jurassic) ammonite succession at Sentralbanken High (core 7533/3-U-1), Barents Sea, and its stratigraphical and palaeobiogeographical significance

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Key words: Middle and Upper Jurassic, ammonites, Boreal and Subboreal provinces, correlation, palaeobiogeography.

Abstract: Ammonites recovered from Upper Bajocian to Upper Kimmeridgian strata in a core drilled at Sentralbanken High in the northern Barents Sea are described and chronostratigraphically interpreted. The lower part of the core comprises Upper Bajocian to Upper Callovian deposits of the Fuglen Formation with ammonites of the genus *Cranocephalites* near the base, and *Longaeviceras* in the upper part. The overlying Oxfordian to Kimmeridgian deposits of the Hekkingen Formation are thin and condensed. The Hekkingen Formation is dominated by Boreal cardioceratid ammonites (*Cardioceras*, *Amoebites*, *Euprionoceras*, *Hoplocardioceras*), with numerous Subboreal aulacostephanid ammonites (*Rasenia*, *Zenostephanus*) at two levels. The occurrences of Boreal and Subboreal ammonites are discussed in relation to the palaeogeography and fluctuations of ammonite faunas within the Boreal Realm during the Kimmeridgian with special attention to levels rich in Subboreal ammonites whose appearance has been controlled by tectonic and climatic factors. The uppermost part of the succession deposited during a time of maximum flooding in the Late Kimmeridgian is dominated by Boreal ammonites.

INTRODUCTION

Shallow stratigraphic drillings on the western Barents Shelf in the past four decades have provided key records of stratigraphic data. Shallow stratigraphic boreholes, generally with continuous cores, were drilled at various localities by SINTEF Petroleum Research (earlier IKU) between 1984 and 1990 and by the Norwegian Petroleum Directorate between 1990 and 2005 (Bugge *et al.*, 1995, 2002; Smelror *et al.*, 2001; Vigran *et al.*, 2014). Ammonite faunas recovered in the cores have proved important tools for high-resolution stratigraphy of the Middle and Upper Jurassic sequences, and for regional correlations between offshore and onshore successions (Wierzbowski, Århus, 1990; Wierzbowski, Smelror, 1993; Wierzbowski *et al.*, 2002).

The present paper concerns ammonites recovered in borehole 7533/3-U-1 drilled in 1990 at the Sentralbanken High in the northwestern part of the Barents Sea (Fig. 1). In addition to a high-resolution biostratigraphic breakdown of the cored succession, the recovered ammonite faunas are used for palaeobiogeographical interpretations. The studied core from the Sentralbanken High comes from a location between the previously investigated sections to the north on Spitsbergen, Kong Karls Land and Franz Josef Land, and to the south in the Nordkapp Basin and offshore Troms III and Nordland VII areas. Through most of the Middle and Upper Jurassic succession which has been studied, the recovered ammonites comprise Boreal Cardioceratidae faunas, but the incoming of Subboreal Aulacostephanidae ammonites at two Kimmeridgian intervals offers possibilities to gain an

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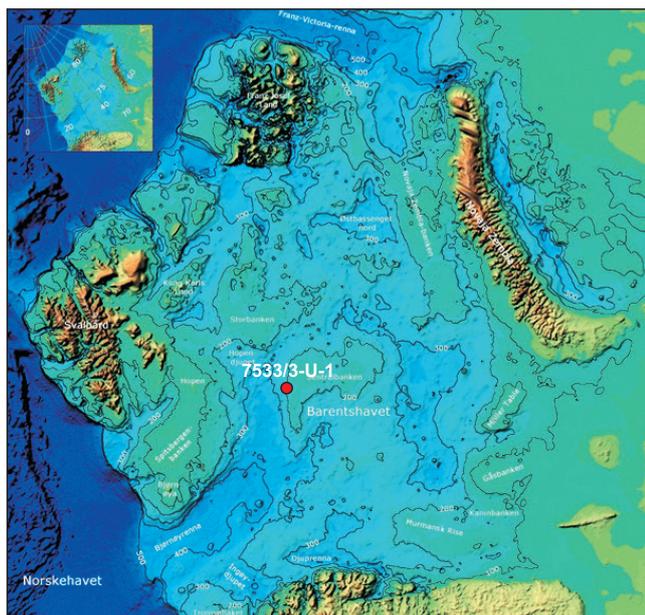


Fig. 1. Location map of borehole 7533/3-U-1 at the Sentralbanken High, Barents Sea

insight into fluctuations in the marine connections between the Boreal and Subboreal provinces during the tectonically active Late Jurassic time.

The ammonites are housed and catalogued in the collection of the Museum of the Geological Faculty of the Warsaw University (collection numbers: MWG UW ZI/98).

CORE DESCRIPTION

Borehole 7533/3-U-1 was drilled through 1.8 m of Quaternary surficial deposits and 118.75 m into Lower Cretaceous and Jurassic bedrocks. The borehole was fully cored, and the core divided into four lithological units (formations), of which the Cretaceous formations are not treated herein. Ammonites which were recovered from the two oldest formations of Middle and Late Jurassic ages, are described (Figs. 2, 3).

The oldest formation, covers the interval from TD at 120.55 up to 71.52 m and consists of dark grey shales and siltstones of Late Boreal Bajocian to Late Callovian (or even earliest Oxfordian) age. The sediments show planar lamination, and intervals with calcareous and sideritic nodules are found throughout the unit. Beds with bioturbation are common in the lower part of the unit, more scattered in the upper part. Belemnites and bivalves were recovered throughout the unit, along with ammonites. The unit was deposited in an open shelf environment and correlates to the Fuglen For-

mation as defined further south in the Barents Sea (Worsley *et al.*, 1988).

The overlying formation from 71.52 to 48.38 m comprises dark grey, organic rich, claystones of Early Oxfordian to Late Kimmeridgian age. Planar lamination is seen through the entire unit. Beds with bioturbation are found in the lower part of the unit (below 66 m). Some belemnites and bivalves are found in the upper part. The unit belongs to the Hekkingen Formation and was deposited in a shelf environment, partly in a setting with restricted bottom-water circulation and anoxic conditions. The boundary with the overlying condensed carbonates of the Early Cretaceous Klippfisk Formation is marked by a stratigraphic break, where Volgian-Ryazanian strata are missing.

AMMONITE STRATIGRAPHY

The two oldest ammonites recovered from 117.24 and 115.67 m in the core belong to the genus *Cranocephalites*. The succession of the *Cranocephalites*-beds has been studied in detail in Jameson Land in East Greenland where a large number of successive faunal horizons were distinguished by Callomon *et al.* (2015). The specimens studied here are compared with the ammonites of that succession. *Cranocephalites* ammonites were also described from northern and north-eastern Siberia by Meledina (1973). However, these ammonites represent a less complete succession of species and the described specimens differ in their ornamentation from those from East Greenland (see Callomon *et al.*, 2015) and the specimens presented herein.

The occurrence of *Cranocephalites indistinctus* at 117.24 m (Pl. 1: 1a–d) is indicative of the Indistinctus Zone, whereas younger *Cranocephalites* sp. at 115.67 m (Pl. 1: 2) may represent any stratigraphical level from the Indistinctus Zone up to the top of the Pompeckji Zone of the so-called *Cranocephalites* beds (Callomon *et al.*, 2015). The ammonites from 117.24 to 115.67 m are indicative of the Boreal Upper Bajocian (Callomon, 2003).

Upwards in the core the ammonite fauna consists of two specimens found at 99.42 and 94.60 m assigned to *Kepplerites* and *Pseudocadoceras*, respectively.

The occurrence of the genus *Kepplerites* at 99.42 m indicates a correlation to the stratigraphical interval from the Boreal Middle Bathonian *Cranocephaloide* Zone and up to the end of the Lower Callovian (Callomon, 1975, 1985, 1993; see also Mitta *et al.*, 2014). Cadoceratids such as the one from 94.60 m, similar to *Pseudocadoceras*, are known to occur from the Boreal Variabilis Zone of the Boreal Upper Bathonian (but some forms transitional to older *Arctico-ceras* occur already in the *Cranocephaloide* Zone; see Cal-

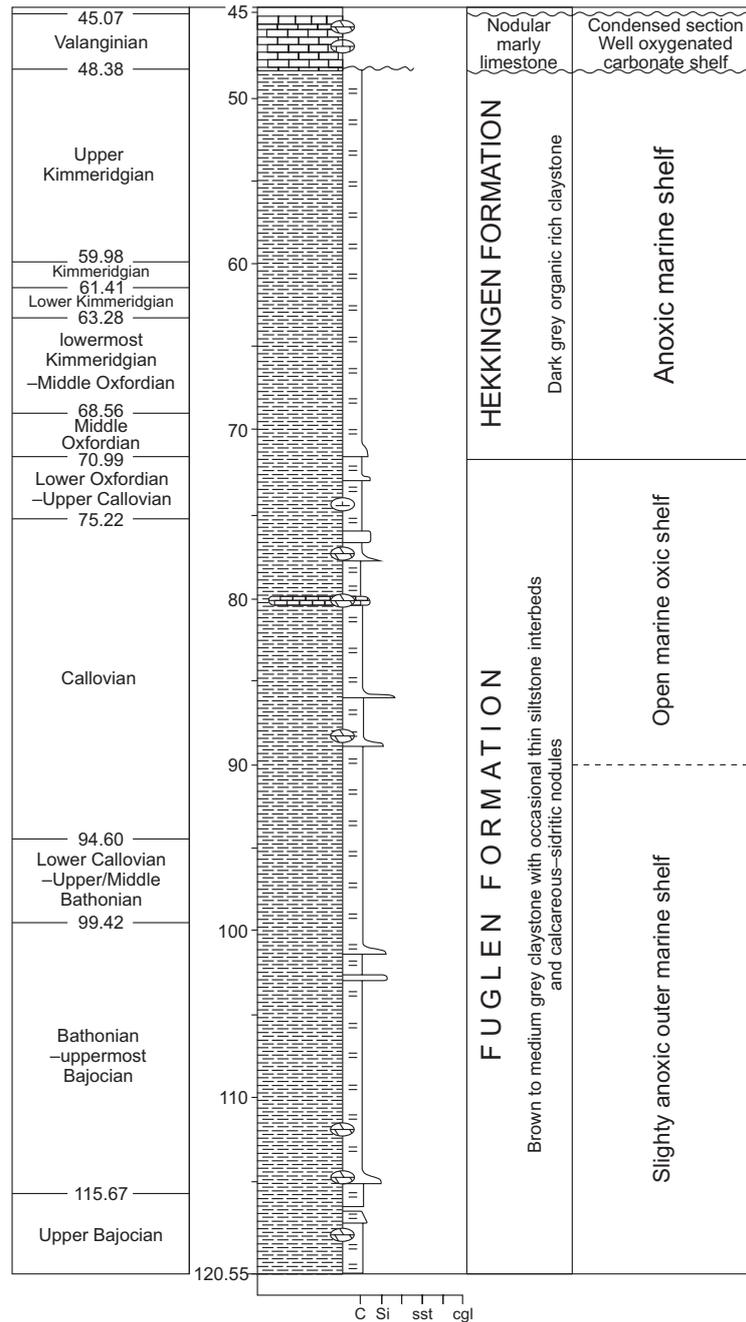


Fig. 2. Lithological log of core 7533/3-U-1 and its stratigraphical and sedimentological interpretation

lomon, 1985), and up to the Middle and even lowest Upper Callovian (Callomon, 1993; Meledina, 1977; Kopik, Wierzbowski, 1988). Taking into account, however, the fact that *Pseudocadoceras* occurs very near to *Keplerites* in the core, as well as the fact that the directly younger ammonite at 75.22 m is indicative of the Upper Callovian (see below), the studied interval between 99.42 and 94.60 m possibly

represents only a fragment of the Middle Bathonian–Lower Callovian succession (Fig. 2).

The ammonite from 75.22 m may be referred to the genus *Longaeviceras*, representing the *L. placenta* (Leckenby) – *L. longaevum* (Leckenby) species (Pl. 1: 3), and is indicative of the Athleta Zone of the Upper Callovian (see Kiselev, Rogov, 2018, and earlier papers cited therein).

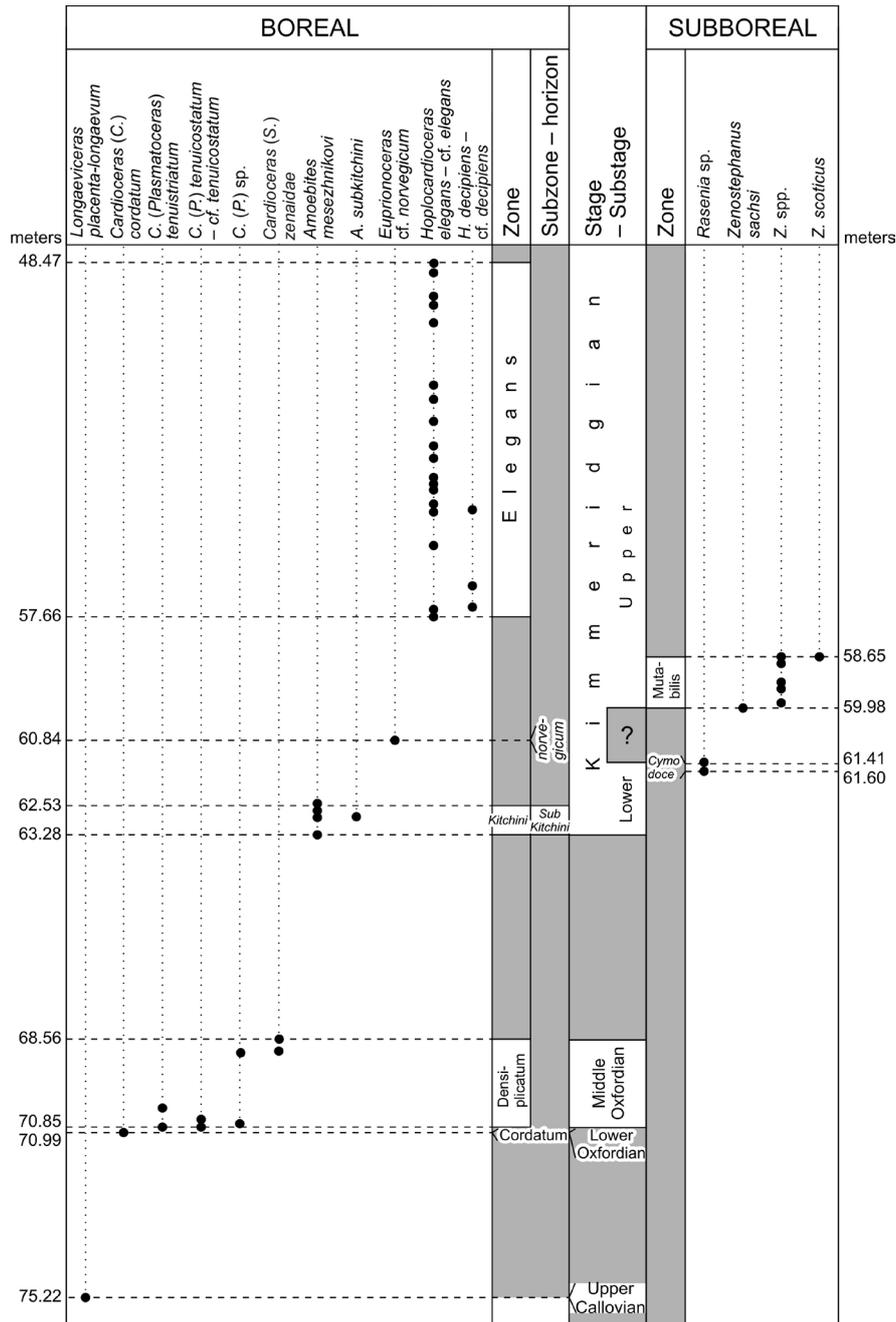


Fig. 3. Ammonite distribution and chronostratigraphical interpretation of the deposits from the Upper Callovian to the Upper Kimmeridgian in core 7533/3-U-1

The overlying younger ammonite faunas, composed entirely of representatives of the genus *Cardioceras*, are typical of the Oxfordian. The specimen found at 70.99 m is *Cardioceras (Cardioceras) cordatum* (Sowerby) (Pl. 1: 4). It indicates the upper part of the Cordatum Zone, i.e. the Cordatum Subzone, which represents the topmost part of the Lower Oxfordian (Arkell, 1941). In accordance with that,

the boundary between the Callovian (as marked by occurrence of *Longaeviceras*) and the Oxfordian, has to be placed in a 4.23 m thick interval, between 75.22 and 70.99 m in the core (Fig. 3).

The overlying ammonite faunas with *Cardioceras* consist of numerous small specimens, showing the fine and dense ribbing typical of the subgenus *Plasmatoceras*. The

following forms are identified: *C. (P.) tenuistriatum* Borissjak (Pl. 1: 5) at 70.85 and 70.37 m, *C. (P.) tenuicostatum* (Nikitin) (Pl. 1: 6) at 70.80 and 70.84 m, and *C. (P.)* sp. at 70.78 and 68.92 m. These ammonites are indicative of a lower part of the Densiplicatum Zone, representing the lowermost part of the Middle Oxfordian (Sykes, Callomon, 1979). The remaining youngest *Cardioceras* at 68.88 and 68.56 m are represented by late forms of the subgenus *Subvertebriceras*, similar to *C. (Subvertebriceras) zenaidae* Ilvoaisky (Pl. 1: 7). They indicate the presence of a higher part of the Densiplicatum Zone of the Middle Oxfordian (Sykes, Callomon, 1979).

All the *Cardioceras* faunas of the Densiplicatum Zone occur from 70.85 to 68.56 m. The boundary between the

Lower and the Middle Oxfordian has to be placed thus in a very narrow 14 cm thick interval between 70.99 and 70.85 m in the core (Fig. 3).

The deposits between 68.56 and 63.28 m did not yield any ammonites. Bivalves occurring in this 5.28 m thick interval (*Entolium*, *Nucula*) are of little stratigraphical importance. The ammonites found directly below and above indicate, however, that the stratigraphical position of the deposits devoid of ammonites may range from the upper part of the Middle Oxfordian (Tenuiserratum Zone) up to the lowermost Lower Kimmeridgian (Bauhini Zone) (see Fig. 4). It is possible that the succession is not complete. The stratigraphical gap may correspond to that which occurs between the Fuglen and the Hekkingen formations in the

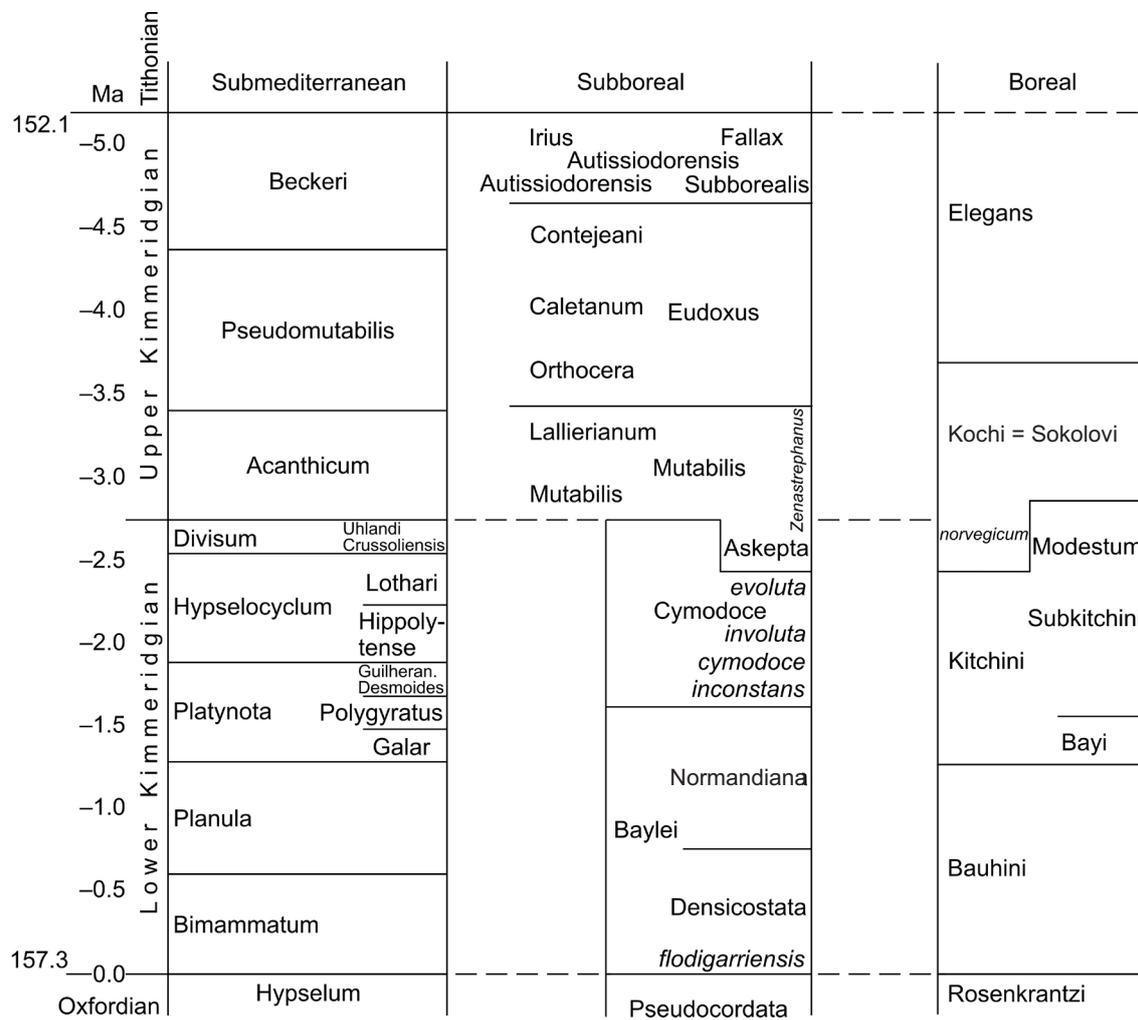


Fig. 4. Correlation of the chronobiostratigraphical ammonite zonation of the Kimmeridgian

The numerical scale of the boundaries of the stage at the left is after Ogg and Hinnov (2012), but the detailed scales are based on biostratigraphic correlations with the succession of the Early Kimmeridgian of south-eastern France where the primary sedimentary cyclicity in the term of short eccentric cycles was recognized by Boullila *et al.* (2010), and of the Late Kimmeridgian in the Kimmeridge Clay in England with eccentricity cycles recognized by (Huang *et al.*, 2010). The details of the biostratigraphical correlations are according to Wierzbowski (2019, and the references given therein)

western part of the Barents Sea, and possibly between the Oppdalen Mbr. and the Lardyfjellet Mbr of the Agardhfjellet Fm. of Svalbard, and embraces either the whole or a part of the stratigraphical interval from the upper Middle Oxfordian to the Upper Oxfordian (Smelror *et al.*, 2001; *cf.* Koevoets *et al.*, 2018, 2019).

Numerous cardioceratids of the genus *Amoebites* are found in the overlying deposits from 63.28 to 62.53 m in the core. One of these is similar to *Amoebites subkitchini* (Spath) at 62.66 m, and the others to *Amoebites mesezhnikovi* (Sykes et Surlyk) (Pl. 1: 8) at 63.28, 62.54 and 62.53 m. The stratigraphical position of the two species is similar – both occur in the Subkitchini Subzone of the middle part of the Boreal Kitchini Zone of the Lower Kimmeridgian (Wierzbowski, Rogov, 2013; Rogov, 2016).

Small aulacostephanids of the genus *Rasenia* (Pl. 2: 1) were encountered at 61.60 and 61.41 m in the core, directly above beds with *A. mesezhnikovi*. They may be compared with early species of the genus *Rasenia* – such as *R. inconstans* (Spath) and/or *R. cymodoce* (d'Orbigny). It is worth noting that the rock samples from the core which yielded *Rasenia* ammonites, as well as the immediately underlying sample from 61.94 m, show the common occurrence of shells of the bivalve *Buchia* [identified as *B. concentrica* (Sowerby)] and the virtual absence of fragments of cardioceratid ammonites. The occurrence of the ammonites of the genus *Rasenia* indicates that the studied interval in the core, possibly from 61.94 m, and surely between 61.60 and 61.41 m devoid of cardioceratids, can be correlated with some lower to middle parts of the Subboreal Cymodoce Zone of the Lower Kimmeridgian. It should be remembered that a similar replacement of the ammonite fauna with *Amoebites mesezhnikovi* by a new one with *Rasenia cymodoce* (with rare *Amoebites*) has been recognized in central and eastern Spitsbergen (Wierzbowski, 1989; Rogov, 2014). The interval indicates a sudden faunal turnover from cardioceratid into aulacostephanid ammonite assemblages, implying a related change in the subdivision of the Kimmeridgian from the Boreal to the Subboreal ammonite zonation (*cf.* Birkelund, Callomon, 1985; Wierzbowski, Smelror, 1993).

The overlying deposits further yielded a cardioceratid fauna at 60.84 m with numerous small ammonites referred to as *Euprionoceras cf. norvegicum* (Wierzbowski) (Pl. 2: 2). These occur together with bivalve *Buchia cf. tenuistriata* (Lahusen). This bivalve also occurs at 60.98 m.

The ammonite *Euprionoceras norvegicum* (Wierzbowski) may be treated as an earliest form of the genus *Euprionoceras* and is indicative of the *E. norvegicum* horizon. The horizon has been recognized as the lowermost part of the Kochi = Sokolovi Zone which was characterized by the occurrence of the ammonite *Euprionoceras* (Wierzbowski, Smelror, 1993). However, the *E. norvegicum* horizon, as de-

finied by the incoming of small densely ribbed *Euprionoceras* ammonites, also correlates with an upper part of the Modestum Subzone, being the uppermost part of the Kitchini Zone (see Fig. 4). Such a correlation results from the co-occurrence of the first representatives of *E. norvegicum* with last representatives of the genus *Amoebites* of the *A. modestum* group in the cores in the southern Barents Sea (*cf.* Wierzbowski, Smelror, 1993). Such a stratigraphical position of the *E. norvegicum* horizon is confirmed by new data from different areas of Europe, showing the co-occurrence or near-occurrence in the sections of the first *Euprionoceras* and the last *Amoebites* (see *e.g.*, Wierzbowski, Rogov, 2013, fig. 4; Wierzbowski *et al.*, 2015, fig. 2; Rogov *et al.*, 2017). This issue is discussed below, and in the chapter on palaeobiogeography.

The succession of the two ammonite groups – the last *Amoebites* and the first *Euprionoceras* as discussed above results in the somewhat nebulous interpretation of the boundary of the two Boreal ammonite zones – the Kitchini Zone, and the Kochi = Sokolovi Zone. The boundary can be placed either at the level where the last *Amoebites* are found, or at the level where the first *Euprionoceras* appear (see Fig. 4, see also Wierzbowski, Smelror, 1993). The latter interpretation is possibly better as it is more univocal and easier for recognition, although the problem needs additional studies. A very similar definition has been accepted recently for the boundary between the Oxfordian and the Kimmeridgian as defined in the Boreal cardioceratid zonation between the Rosenkrantzi Zone (uppermost Oxfordian) and the Bauhini Zone (lowermost Kimmeridgian). The base of the Bauhini Zone has been placed at the level where the first ammonites of the genus *Plasmatites* appear, which co-occur here with the last *Amoeboceras* ranging up from the underlying Rosenkrantzi Zone (Wierzbowski *et al.*, 2018).

A younger fauna occurs from 59.98 to 58.65 m and is characterized by the incoming of aulacostephanid ammonites, represented by numerous specimens of *Zenostephanus* (formerly *Xenostephanus*). The specimen at 59.98 m (Pl. 2: 3) is very close to *Zenostephanus sachsi* (Mesezhnikov) (see Mesezhnikov, 1969a, p. 119–120, pl. 17: 1; pl. 19: 1; pl. 20: 1). The specimens found at 59.87, 59.40, 59.37, 58.88 and 58.71 m are too fragmentarily preserved to be specifically determined, although all of them belong to the genus *Zenostephanus*. The youngest specimen of *Zenostephanus* referred to as *Z. scoticus* (Arkell et Callomon) (Pl. 2: 4) is found at 58.65 m in the core.

Hence, the deposits between 59.98 to 58.65 m represent the *Zenostephanus*-rich level, which is a typical aulacostephanid horizon (*Zenostephanus* horizon) recognized widely in the Boreal Realm. This level may be correlated mostly with the Subboreal Mutabilis Zone (especially its part distinguished often as the Mutabilis Subzone) (see Ro-

gov, Poulton, 2015, and references given therein; see also Fig. 4).

The youngest faunal assemblage in core 7533/3-U-1 occurs between 58.46 m (*i.e.* directly above the youngest recovered *Zenostephanus*) and it continues up to 48.47 m. It is composed entirely of cardioceratid ammonites. The two specimens at 58.46 and 58.15 m are very poorly preserved and cannot be interpreted unequivocally even to the genus/subgenus level. The rest of the very numerous (over 20) specimens from 57.66 to 48.47 m (in about a 9.20 m thick interval) belong to the genus *Hoplocardioceras* which is indicative of the Boreal Elegans Zone as interpreted by Wierzbowski and Smelror (1993). The genus includes two closely related species or subspecies – *H. elegans* (Spath) (Pl. 2: 5, 7, 8) and *H. decipiens* (Spath) (Pl. 2: 6), which show “at least marked overlapping if not total covering of their stratigraphic ranges which precludes the differentiation of smaller biostratigraphic units of wider importance within the Elegans Zone” (Wierzbowski, Smelror, 1993, p. 246). However, some local differences in the stratigraphical ranges of the two forms have been observed: *H. elegans* dominates in dark micaceous shales, and *H. decipiens* often occurs in the proximity of lumachelle beds (*cf.* Wierzbowski, Smelror, 1993). In relation to the core studied – *H. decipiens* occurs together with *H. elegans* in the lowermost part of the stratigraphical range of the genus *Hoplocardioceras* (from 57.66 to 54.84 m), whereas *H. elegans* without *H. decipiens* occurs in its middle and upper parts (from 54.72 to 48.47 m).

CHANGES IN AMMONITE FAUNAS AS A CLUE FOR PALAEOBIOGEOGRAPHICAL INTERPRETATION

The stratigraphical interval from the Late Bajocian to the end of the Kimmeridgian in the studied core shows the presence of Boreal and Subboreal ammonites. All these ammonites were typical of the Boreal Realm which stretched over wide areas of high latitudes of the northern hemisphere (*e.g.*, Cariou *et al.*, 1985; Callomon, 2003; and earlier papers cited therein). It is commonly accepted that the Boreal ammonites, represented by the successive members of the family Cardioceratidae, developed continuously from the Late Bajocian to the end of the Kimmeridgian in parts of the Boreal Sea – a more or less enclosed circum-polar marine basin (Callomon, 1985). On the other hand, a peripheral area of the Boreal Realm was inhabited by ammonite groups of Subboreal affinities (*e.g.*, Cariou, 1973; Matyja, Wierzbowski, 1995). These included the family Kosmoceratidae from the Middle Bathonian to the end of the Callovian, as well as the family Aulacostephanidae from the Late Oxfordian to the end of the Kimmeridgian (*e.g.*, Birkelund *et al.*, 1983;

Mesezhnikov, 1984; Cariou *et al.*, 1985; Birkelund, Callomon, 1985; Matyja *et al.*, 2006; Wright, 2010).

The distribution, and relations between the Boreal and Subboreal ammonite faunas during the Middle Jurassic is not discussed herein. This because of the poor recovery of these faunas in the present core. During the latest Callovian and Early Oxfordian to Middle Oxfordian separate Subboreal ammonite faunas as defined above did not exist due to a wide marine transgression and the opening of the Boreal Realm towards the south. Thus, the comments given below are mostly related to the palaeogeographic distribution of the ammonite faunas during the Late Oxfordian and the Kimmeridgian.

In contrast the more variegated and fluctuating environmental conditions in the Subboreal Province, the conditions of the Boreal Province were more uniform, although even here, especially during Kimmeridgian, occurred areas and levels with shales showing nothing but *Buchia* bivalves, and possibly deposited in an anoxic environment. In these deposits cardioceratid ammonites “occur at isolated, well-spaced but widespread horizons, suggesting infrequent and highly episodic colonization of an otherwise hostile environment” (Callomon, 1985, p. 73). The Subboreal Province was characterized by a diversified and strongly contrasted facies pattern generally related to the existence of a more shallow-water environment, compared to the deeper-water conditions in the Boreal Province (Fürsich, Sykes, 1977).

Although the Subboreal Province was originally referred to a relatively small area at the southern rim of the Boreal Realm, *i.e.* from North-West Europe to northern areas of the European part of Russia, west of the Ural Mountains (*e.g.*, Sykes, Callomon, 1979), it is becoming evident that some other areas placed in more inner parts of the Boreal Realm should also be included in the Subboreal Province because of the common occurrence of the diagnostic aulacostephanids, at least in some time intervals. The intricate relations between the areas with Boreal and Subboreal ammonite faunas were first recognized in northern Central Siberia. Apparently an east-west stretching deep marine basin characterized by the occurrence of Boreal cardioceratid ammonites was developed during the Late Oxfordian to Kimmeridgian between two shallower sea areas bordering Taimyr Island to the north, and Middle Siberian Land to the south (Mesezhnikov, 1969a, b; see also Zakharov, Shurygin, 1983; Zakharov *et al.*, 2005). This area was recognized by Mesezhnikov (1969b) as corresponding to the Taimyr-Canadian Province, coeval with the Boreal Province as interpreted herein. On the other hand, the shallower areas to north and south, bordering the land-masses mentioned were inhabited by numerous aulacostephanids typical of the Subboreal Province, and distinguished as the Greenland-Khatanga Province by Mesezhnikov (1969b).

Another example of a specific distribution of the Boreal and Subboreal faunas is related to the existence of a deep-water marine basin of the West Siberian Sea stretching over vast areas of Western Siberia. Here the most common ammonite occurrences in the Kimmeridgian “belong to ammonite subgenus *Amoeboceras* (*Amoebites*)” (Alfirov *et al.*, 2016) – *i.e.* typical faunas of the Boreal Province, although Subboreal aulacostephanids appear at some levels.

When discussing the palaeogeographic pattern of the Barents Sea area during Late Oxfordian and Kimmeridgian times, special attention should be paid to the ammonite faunas of the Nordkapp Basin (Fig. 5). During the Late Jurassic the around 300 km long basin in the southwestern Barents Sea represented a deep marine embayment founded on top of a Late Palaeozoic rift basin (Bugge *et al.*, 2002). The ammonite faunas of the Late Oxfordian and Kimmeridgian studied in detail in several cores (Wierzbowski, Århus,

1990; Wierzbowski, Smelror, 1993) are here composed exclusively of different cardioceratids typical of the Boreal Province.

Core 7533/3-U-1 is from Sentralbanken High located north of the Nordkapp Basin, closer to the Svalbard Archipelago (Figs. 1, 5). The ammonite faunas consist mostly of Boreal Cardioceratidae from the Upper Bajocian up to the Middle Oxfordian and the lower part of the Lower Kimmeridgian (from 117.24 to 62.53 m). In between there is an interval devoid of ammonites.

In borehole 7533/3-U-1 an abrupt appearance of the Subboreal ammonites of the genus *Rasenia* indicative of the lower to middle parts of the Lower Kimmeridgian Cymodoce Zone is observed at 61.60 to 61.40 m in the core (Figs. 3, 4). It is evidently related to changes in environmental conditions. The same phenomenon of a sudden appearance of *Rasenia* ammonites, after a long time of existence of Bo-

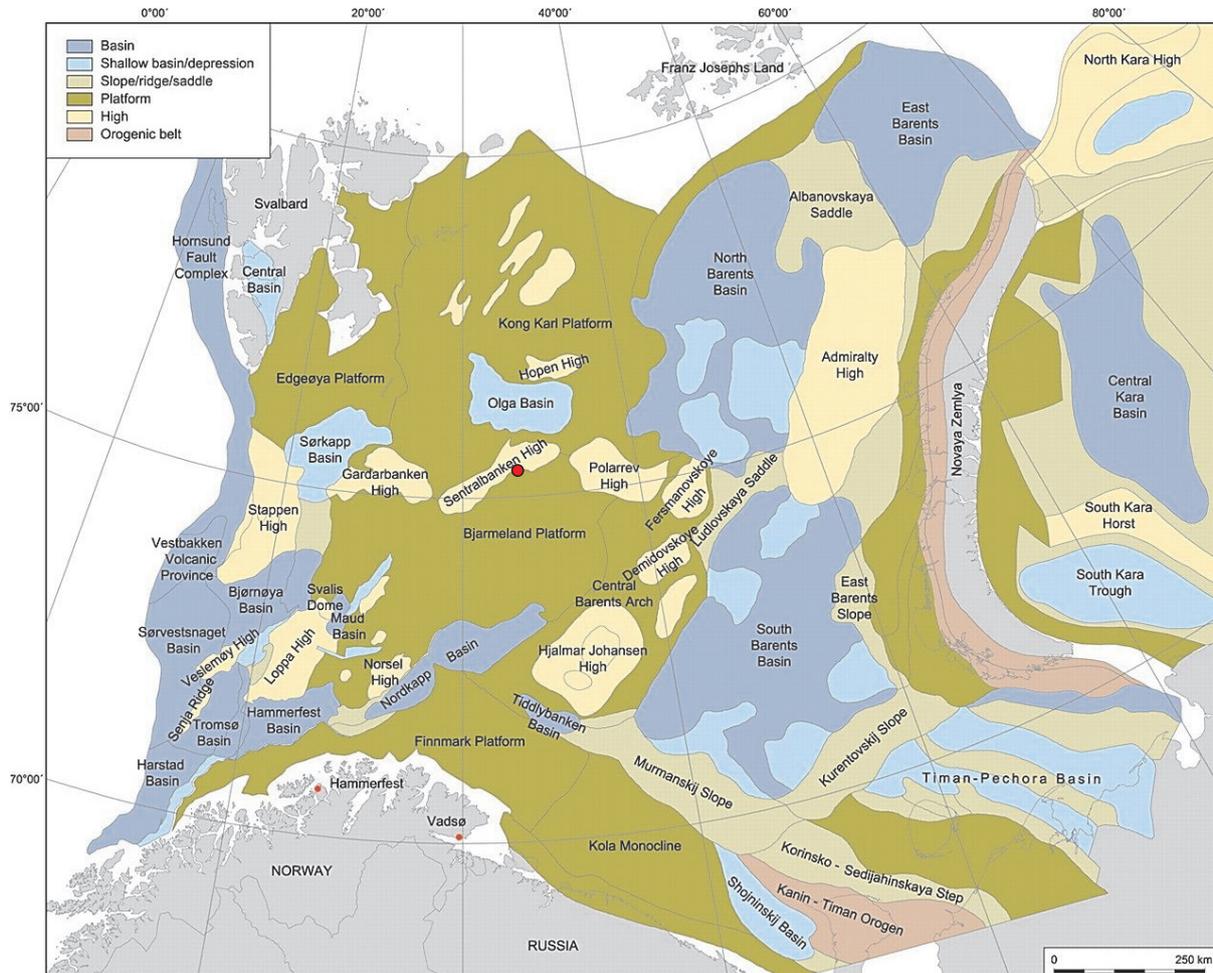


Fig. 5. Geological map showing main structural elements of the Barents Sea with location of borehole 7533/3-U-1 (in red) at the Sentralbanken High (after Henriksen *et al.*, 2011)

real cardioceratids, was recognized in central and eastern Spitsbergen (Wierzbowski, 1989; Rogov, 2014). It is also recognized in Andøya Island (Lofoten, northern Norway) where the *Rasenia* ammonites of the Cymodoce Zone appears in the Breisanden Member of the Dragneset Formation. These are the first ammonites occurring here above the disconformity between the Bajocian fluvial sediments below, and the shallow-marine deposits above (Birkelund *et al.*, 1978; Smelror *et al.*, 2001; with earlier references cited therein). Another common occurrence of *Rasenia* ammonites was initiated within the deep shelf sediments of the Alge Member of the Hekkingen Formation in the Nordland VII offshore area in the northeastern part of the Norwegian Sea (Wierzbowski *et al.*, 2002). Much further south, outside the Boreal Realm, a sudden occurrence of the ammonite *Rasenia cymodoce* (d'Orbigny) is observed within the Submediterranean ammonite succession of Aquitaine, western France (Hantzpergue, 1989 and his earlier papers cited therein). This visit of Subboreal ammonites is approximately coeval with the *Rasenia* occurrences discussed above, and it indicates "the abrupt arrival which reflects an invasion from the north" (Birkelund *et al.*, 1983, p. 305).

All these observations show that the Subboreal *Rasenia* ammonites appeared within a relatively short time interval in the Early Kimmeridgian, and over widespread areas. Such a migration appears to be partly tectonically induced or related to an overall transgressive phase and may correspond to rifting within the Boreal Sea, both in European and Arctic areas. The tectonic movements and relative sea-level rise resulted in flooding of previous elevated areas or shallow marine areas and deepening of shelves, such as Andøya in Northern Norway, the offshore Nordland VI and Troms III areas, and on East Greenland (Birkelund *et al.* 1978; Surlyk, 1978; Smelror *et al.*, 2001; Wierzbowski *et al.*, 2002), with possible maximum sea-level being reached in the mid Kimmeridgian (Surlyk, 1990). The overall transgression and subsidence of marine barriers would have opened new sea-routes for migration and enabled the settlement and flourishing of the Subboreal ammonite faunas of *Rasenia*. Increased sea-water temperatures may also have played a role (*e.g.*, Zakharov *et al.*, 2005; Rogov, 2014), by enabling Subboreal ammonites to migrate into such a remote area of the Boreal Sea as the Svalbard and the Franz Josef Land archipelagos previously devoid of Aulacostephanidae (see *e.g.*, Mesezhnikov, Shulgina, 1982; Mesezhnikov, 1984; Wierzbowski, 1989). This may explain the long migration route of the genus *Rasenia* (close to *R. cymodoce*) recently described from the Pacific Coast of western Canada in British Columbia (Rogov, Poulton, 2015).

The sudden appearance of the Boreal cardioceratids (*Euprionoceras cf. norvegicum*) at 60.84 m in the core has a temporarily character (Fig. 3). Immediately above, from

59.98 to 58.65 m, once more the Subboreal aulacostephanids (*Zenostephanus*) occur. Nevertheless, it is important for stratigraphical and biogeographical interpretations. The occurrence of the completely new ammonite faunas composed of small-sized *Euprionoceras*; associated in some cores from the Barents Sea with the last *Amoebites* (see Wierzbowski, Smelror, 1993), as well as the appearance directly above of a new fauna of *Zenostephanus*, marks a major reorganization of the whole ammonite faunas of the Boreal Realm. This may be correlated with a large transgression which occurred over wide areas of the northern Tethyan Shelf near the boundary between the Early and Late Kimmeridgian, and which changed markedly the distribution and composition of the ammonite faunas in large areas of Europe (Matyja, Wierzbowski, 2000, see also Fig. 4).

The transgression is well known in the southern part of the Boreal Realm and the adjoining areas transitional to the Submediterranean Province. Here it resulted in a wide distribution of the aulacostephanid *Rasenioides*, gradually followed phylogenetically by *Aulacostephanoides* in beds distinguished as the Askepta Subzone (Birkelund *et al.*, 1978, 1983). The Askepta Subzone is correlated with the upper part of the Cymodoce Zone in the Subboreal zonal scheme in NW Europe (*i.e.* the top of the Lower Kimmeridgian), or it is treated as a lower part of the Mutabilis Zone (*i.e.* the lowest part of the Upper Kimmeridgian) (Birkelund *et al.*, 1978, 1983; Hantzpergue, 1989, 1995; Matyja, Wierzbowski, 2000; see also Fig. 4 herein). In this context the boundary between the Subboreal Cymodoce Zone and the Mutabilis Zone appears particularly vague as these zones are based on ammonites representing the two different aulacostephanid branches, separated not only phylogenetically, but also provincially (Birkelund *et al.*, 1983). The Cymodoce Zone is based on the stratigraphical range of the genus *Rasenia*, whose natural descendent is the genus *Zenostephanus*. The Mutabilis Zone is based on the genus *Aulacostephanoides*, which natural forerunner is the genus *Rasenioides* (see *e.g.*, Arkell, Callomon, 1963; Birkelund *et al.*, 1983; Birkelund, Callomon, 1985).

Recent biostratigraphical studies in Tatarstan (central European Russia) have revealed ammonite assemblages with *Rasenioides* and *Aulacostephanoides* together with small cardioceratids of the genus *Amoebites*, and some forms transitional to the genus *Euprionoceras*. These occur along with the last raseniids (*Eurasenia* or *Involuticeras*), and directly below *Zenostephanus*, in the stratigraphical interval at the top of the Cymodoce Zone (Askepta Subzone), and a lower part of the Mutabilis Subzone of the Mutabilis Zone (Rogov *et al.*, 2017, fig. 4). This, as well as observations from the cores in the Nordkapp Basin (Wierzbowski, Smelror, 1993), suggest that the older genus *Amoebites* was gradually replaced by the younger genus *Euprionoceras* at

the transition between the Cymodoce Zone and the Mutabilis Zone *sensu stricto*, i.e. in the Askepta Subzone. It seems thus highly probable that the deposits at 60.84 m in core 7533/3-U-1, which yielded the oldest small-sized *Euprionoceras* referred to as *E. cf. norvegicum*, may correspond at least partly to the stratigraphical interval of the Askepta Subzone (Figs. 3, 4).

The ammonites of the genus *Zenostephanus* from 59.98 to 58.65 m in the core were distinguished for the first time under the name “*Xenostephanus* (M) – *Xenostephanoides* (m)” from Kimmeridgian sandy deposits of local origin in Lincolnshire, north-eastern England discovered in glacial drift (Arkell, Callomon, 1963). Two aulacostephanid assemblages were recognized therein in deposits of the same lithology, partly together in the same Quaternary locality, but partly isolated in different localities: (A) composed of *Rasenia*, *Rasenioides* and *Aulacostephanoides*; (B) composed of “*Xenostephanus* – *Xenostephanoides*” (i.e. of *Zenostephanus* as interpreted currently). Additionally, there were recognized small-sized cardioceratids of the genus *Amoebites* such as *A. kitchini* (see Arkell, Callomon, 1963, pl. 32: 26a, b) and a form described as *A. cricki* or cf. *cricki* (Arkell, Callomon, 1963, pl. 32: 25a–c) – which looks in fact very similar to early *Euprionoceras*. All these ammonites were treated either as of “the same age which is Mutabilis Zone *sensu stricto* (...)”, or if not strictly coeval, “the differences of age (...) could not have been large” (Arkell, Callomon, 1963, p. 240–241).

The genus *Zenostephanus* is widely recognized in several Arctic areas like Franz Josef Land (Mesezhnikov, Shulgina, 1982; Repin *et al.*, 2007), and Spitsbergen (Friebold, 1930: see Wierzbowski, 1989; Rogov, 2014; Rogov, Poulton, 2015, and earlier papers cited therein; Koevoets *et al.*, 2018). It is also known from the Pacific Coast of western Canada (British Columbia), which occurrence possibly was a result of expansion of the Arctic areas (Rogov, Poulton, 2015). In all these areas, ammonites of the genus *Rasenia* were also reported, but their close phylogenetic relation was not recognized in detail. The only stratigraphical interval and area showing the continuous phylogenetical transition between the genera *Rasenia*–*Zenostephanus* through forms of *Zonovia* type, is the Uralensis Subzone of the Evoluta Zone in Northern Siberia at the Kheta River Basin which corresponds approximately to the Askepta Subzone (Mesezhnikov, 1969a, 1984; Saks *et al.*, 1969a, b; see also Wierzbowski, Rogov, 2013). Here, the overlying deposits yielded also *Aulacostephanoides mutabilis* (Sowerby) together with *Zenostephanus*, corresponding to the Mutabilis Subzone. This observation strongly suggests that the occurrence of the genus *Zenostephanus* without *Rasenia*, as observed in core 7533/3-U-1 between 59.98 to 58.65 m, represents the Mutabilis Subzone well above the Askepta Subzone.

The youngest deposits from 58.46 to 48.47 m in the studied core yielded only numerous cardioceratids. The oldest ones (58.46 to 58.15 m) are poorly identifiable, but all the younger specimens belong to the genus *Hoplocardioceras*. This genus is indicative of the Boreal Elegans Zone of the Upper Kimmeridgian, which corresponds to the bulk of the Subboreal Eudoxus Zone, and at least a part of the Autisiodorensis Zone (Wierzbowski, Smelror, 1993; cf. Birkelund, Callomon, 1985; see also Fig. 4 herein). The wide occurrence of these cardioceratids indicates the transgression which occurred over large areas of Europe, including the Boreal Realm, during the early Eudoxus Chron (e.g., Birkelund *et al.*, 1983; Surlyk, 1990; Hantzpergue, 1995; Gallois, 2016). This transgression flooded shallow areas in the Arctic inhabited previously by aulacostephanids, including the western Barents Sea Shelf (Smelror *et al.*, 2001). This possibly resulted in the sudden appearance of the last members of the Aulacostephanidae lineage, such as *Aulacostephanus/Aulacostephanites*, much towards the south from the circum-polar Arctic areas, and down into more southern, and shallow parts of the Boreal Realm and adjoining parts of the Submediterranean Province (see e.g., Birkelund *et al.*, 1985).

NOTES ON AMMONITE IDENTIFICATIONS AND SYSTEMATICS

Below some notes on the identification and systematics of selected species and specimens are added. The *Cranoccephalites* from 117.24 m (Pl. 1: 1a–d) comprises the phragmocone up to about 60 mm diameter, and the body-chamber about 0.75 of a whorl long (but without its outer part – cut off during drilling). However, the whole specimen reached about 70 mm in diameter. The ribbing is uniformly fairly dense in the phragmocone (about 20 primary ribs per whorl), consisting of similarly developed primary and secondary ribs, but the ribs become more widely spaced and coarser at the end of phragmocone – beginning of the body-chamber, and the ribbing finally disappears at the end of body-chamber. These features show that the specimen may be referred to *Cranoccephalites indistinctus* Callomon, showing especially large similarity to the form recognized as *Cranoccephalites indistinctus* trans α' of Callomon *et al.* (2015, pl. 5: 1–3). Another ammonite from 115.67 m is a fragment of the inner whorls of a young specimen possibly about 25 mm in diameter (Pl. 1: 2). The ribbing consists of biplacate ribs with some intercalatory ribs in the outer part of whorl (the secondary/primary ribs ratio equals about 2.5; the primary ribs are strongly prorsiradiate, and the secondaries show a rectiradiate course). This character of ribbing is typical of the inner whorls of several species of *Cranoccephalites*

– except the oldest one – *C. borealis* (Spath), all of them showing similarly developed primary and secondary ribs (see Callomon, 1993; Callomon *et al.*, 2015).

The specimen found at 99.42 m is assigned to *Kepplerites*. This is about 25 mm in diameter, shows evolute coiling, and has short primary ribs which become more distant on the outer whorl preserved. The specimen resembles very much the inner whorls of *Kepplerites*, and it can be safely referred that genus. However, a more detailed determination is difficult. The specimen from 94.60 m is also small, about 25 mm in diameter, shows involute coiling, markedly prorsiradiate strong ribs, and very deep umbilicus. The specimen is undoubtedly a cadoceratid, and may be attributed to *Pseudocadoceras* as it shows the microconch morphology of the cadoceratoid type.

The ammonite from 75.22 m may be referred to the genus *Longaeviceras*. It is large and heavily flattened specimen, showing a narrow umbilicus, and moderately dense ribbing with a fairly low and irregular point of ribs' division (Pl. 1: 3). It is close to late representatives of the genus such as *L. placenta* (Leckenby) and *L. longaevum* (Leckenby), which show a similar morphology, and rather minor differences, doubtfully “more than varietal” in character (Callomon, Wright, 1989).

The ammonites of the genus *Cardioceras* include the type species – *C. (C.) cordatum* (Sowerby) at 70.99 m (Pl. 1: 4) with the primaries well differentiated and loosely-placed, and the numerous, well separated and strongly accentuated secondaries. The finely ribbed subgenus *Plasmaticeras* is represented by specimens belonging to two species: (1) a more distinctly ribbed *C. (P.) tenuicostatum* (Nikitin), and (2) the weakly ribbed *C. (P.) tenuistriatum* Borissiak, the latter includes also a very densely ribbed specimen at 70.37 m (Pl. 1: 5), somewhat resembling in its very delicate ribbing a younger *Mitiscardioceras*. The youngest specimen of the genus *Cardioceras* at 68.56 m (Pl. 1: 7) shows strongly developed ribbing consisting of loosely-spaced primaries, and short secondaries, as well as a coarsely serrated keel. These are typical features of the subgenus *Subvertebriceras* and the specimen can be attributed to *Cardioceras (Subvertebriceras) zenaidae* Ilovaisky. The discontinuity between the primary and secondary ribs along with the presence of strongly developed mid-lateral and ventrolateral tubercles observed at the end of the last whorl preserved, representing already the beginning of the body-chamber, approaches somewhat *Maltoniceras* as well as *Cawtoniceras*.

Numerous cardioceratids of the genus *Amoebites* are found in the overlying deposits from 63.28 to 62.53 m in the core. Two forms may be distinguished based mostly on fragmentarily preserved specimens (some of a larger size with whorl-height attaining even 30 mm): one similar to *Amoebites subkitchini* (Spath) at 62.66 m, and another to *Amoebites mesezhnikovi* (Sykes et Surlyk) (Pl. 1: 8) at 63.28, 62.54, and 62.53 m. These two species are closely related, and differ mostly in the number of secondary ribs (or corresponding ventrolateral tubercles at growing diameters), which is larger in *A. mesezhnikovi* (see Wierzbowski, 1989; see also Sykes, Surlyk, 1976).

The small aulacostephanids of the genus *Rasenia* (Pl. 2: 1) encountered at 61.60 and 61.41 m. show weakly evolute coiling, short primary ribs (26 primary ribs per whorl), and about 2–3 secondaries per each primary rib at about 35 mm diameter. The constrictions are fairly prominent, bordered by stronger ribs. The specimens are too small for unequivocal specific determination. However, the character of ornamentation suggests that they may be compared with early species of the genus *Rasenia* – such as *R. inconstans* (Spath) and/or *R. cymodoce* (d'Orbigny).

The cardioceratid fauna found at 60.84 m is composed of numerous small specimens referred to as *Euprionoceras cf. norvegicum* (Wierzbowski). They are about 12–15 mm in diameter. All of them show the inner whorls up to about 10 mm in diameter, almost smooth or covered by thin striae, followed by biplicate ribs splitting high on the whorl side, and additionally some intercalatory ribs. All these specimens, although poorly preserved, may be referred to as *Euprionoceras cf. norvegicum* (Wierzbowski) (Pl. 2: 2). Such small-sized forms showing dense and rather weak ribbing without markedly developed tubercles represent a fragment of the cardioceratid lineage preceding the appearance of large forms of the genus *Euprionoceras* (see Wierzbowski, Rogov, 2013, fig. 4; Wierzbowski, Smelror, 1993).

The ammonite fauna from 59.98 to 58.65 m is represented by numerous specimens of *Zenostephanus* (formerly *Xenostephanus*) which show characteristic ornamentation and coiling typical of the genus, but as the specimens are strongly flattened they do not reveal the ventral side of the whorl with the characteristic smooth band. The oldest specimen at 59.98 m (Pl. 2: 3) is very close to *Zenostephanus sachtsi* (Mesezhnikov) (see Mesezhnikov, 1969a, p. 119–120, pl. 17: 1; pl. 19: 1; pl. 20: 1). It is about 70–80 mm in diameter (not complete), shows about 12–13 primary ribs per half a whorl at 30–70 mm diameter, bi- and triplicate ribs on the last whorl, and a constriction bordered in its front by a single rib. Several younger specimens found at 59.87, 59.40, 59.37, 58.88 and 58.71 m are too fragmentarily preserved to be specifically determined, although all of them belong to the genus *Zenostephanus*.

The youngest specimen of *Zenostephanus* found at 58.65 m in the core is well preserved (Pl. 2: 4). The specimen attains about 70 mm in diameter as calculated from the preserved fragment in the core. It consists of a phragmocone up to about 65 mm diameter, and a fragment of the body-

chamber. Its inner whorls up to 15 mm diameter show triplicate ribs of the *Rasenia* type, but soon thereafter biplicate ribs occur commonly, covering the next three whorls; the ribbing consists of strong short primaries separated from secondaries by a smooth band – which is typical feature of *Zenostephanus*. The ribbing is dense (from about 24 primary ribs per whorl at 20 mm diameter up to 33 at 50 mm diameter). More loosely spaced triplicate ribs appear on the part of the outer whorl preserved. The coiling is very evolute (at 52 mm diameter, the whorl height and the umbilical width represent 28%, and 47.1% of the diameter, respectively). The specimen compares well with several macroconchs of *Zenostephanus* from Franz Josef Land described under the name *Zonovia* aff. *sachsi* Mesezhnikov by Mesezhnikov, Shulgina (1982, p. 28, fig. 3a, b) and *Zonovia* (*Xenostephanus*) *sachsi* Mesezhnikov by Repin *et al.* (2007, pl. 12: 2a, b; pl. 14: 1, 2a, b, 4). However, all these specimens differ from typical specimens of *Z. sachsi* in several features, such as a markedly denser ribbing and the dominant biplicate ribs in the middle whorls (*cf.* Mesezhnikov, Shulgina, 1982; see also Mesezhnikov, 1984). Very similar specimens referred to as “*Rasenia* sp. indet. *cf. groenlandicus* Ravn” have been described by Frebold (1930, p. 62, pl. 22: 2) from Spitsbergen and these “may well be more closely related to *Xenostephanus*” (*i.e.* *Zenostephanus* as currently interpreted) (Birkelund, Callomon, 1985, p. 39–40). The form represented by the discussed specimens thus deserves to be distinguished as a separate species.

There is no doubt that the discussed specimens, showing the common occurrence of biplicate ribs in the middle whorls which are rather densely-placed, represent another species of the genus *Zenostephanus* than the more heavily ornamented species like *Z. ranbyensis* (Arkell et Callomon) or *Z. sachsi* (Mesezhnikov). It should be also remembered that the discussed specimens show marked similarity to such forms as “*Rasenia borealis* Spath (see Spath, 1935, pl. 6: 1; pl. 7; and Birkelund and Callomon, 1985, pl. 15: 2; pl. 21: 2, 3) but also “*Aulacostephanus groenlandicus* Ravn (see Birkelund, Callomon, pl. 21: 1a, b; see also Sykes and Surlyk, 1976, fig. 7A) which are, however, less evolute and show less swollen primary ribs. The only species which is close to the discussed specimens of *Zenostephanus* is *Zenostephanus scoticus* (Arkell et Callomon), as based on a fragmentarily preserved aulacostephanid ammonite from Scotland, illustrated by Spath (1935, p. 49, pl. 13: 8a, b). It was originally recognized as a microconch of *Xenostephanus* = *Zenostephanus*, and thus placed in the subgenus *Xenostephanoides* by Arkell and Callomon (1963), but such a dimorphic interpretation is not well settled. The name *Zenostephanus scoticus* (Arkell et Callomon) when used both for macro- and microconchs may be thus introduced for the discussed specimen from the present core (Pl. 2: 4).

The moderately to heavily ornamented ammonites recognized at the top of the succession of the core from 57.66 to 48.47 m show commonly the presence of the ventrolateral nodes and clavi typical of *Hoplocardioceras elegans* (Spath) (Pl. 2: 5, 8). The largest of them are represented as a part of the final whorl with ventral rostrum, attaining even about 40–45 mm in whorl height (specimens from 54.22 and 49.49 m), which indicates that the whole specimen may have reached even about 140 mm in diameter (*cf.* Wierzbowski, 1989, pl. 21: 3). Some of the specimens like that from 54.84 m (Pl. 2: 7) show more sparsely-placed ribbing at a diameter about 25 mm and strongly accentuated nodes in the ventrolateral and middle part of the whorl, resembling somewhat the three-tuberculate *H. decipiens* (Spath). The small specimen from 56.83 m (Pl. 2: 6) showing three well developed rows of tubercles, already at about 20 mm diameter and sparse ribbing (about 25 ribs per whorl), is attributed to *Hoplocardioceras decipiens* (Spath).

CONCLUSIONS

The Jurassic succession in core 7533/3-U-1 from the Sentralbanken High of the Barents Sea can lithostratigraphically be subdivided into the Fuglen and Hekkingen Formations with their boundary at 71.52 m, and biostratigraphically into three parts showing different ammonite assemblages and sedimentary patterns: the lowest corresponding to the Middle Jurassic (117.24 to 75.22 m), the middle corresponding to the Oxfordian and Lower to lowermost part of the Upper Kimmeridgian (70.99–58.65 m), and the upper corresponding mostly to the Upper Kimmeridgian Elegans Zone (and the Subboreal Eudoxus Zone) from 58.46 to 48.47 m (Figs. 2, 3). These can be compared with coeval, stratigraphically well described deposits of Spitsbergen and elsewhere in the Arctic.

The oldest part includes nearly the whole Middle Boreal Jurassic from the Upper Bajocian to the Upper Callovian and is fairly complete (*i.e.* defined from the base, middle and top by ammonites and not showing any sedimentological features of discontinuities), attaining about 43 metres in thickness. It is stratigraphically more complete than the corresponding Middle Jurassic deposits of the Svalbard Archipelago. The latter as known from Spitsbergen are generally thinner, although showing some local differences in thicknesses and possibly in stratigraphical range of the deposits. The deposits are possibly not older than the Middle Boreal Bathonian (Cranocéphaloide Zone), as shown by the presence of ammonites of the genus *Keplerites* in the central part of Spitsbergen – in the Sassenfjorden area (Kopik, Wierzbowski, 1988, and earlier papers cited therein), but locally also older – of the Early Bathonian age (Ishmae Zone),

as proved by occurrence of *Arcticoceras*, in eastern and southern Spitsbergen (*cf.* Ershova, 1983). A similar assemblage of ammonites of Early Bathonian age composed mostly of *Arcticoceras* from the Oppdalen Member, in the lowermost part of the transgressive deposits of the Agardhfjellet Formation (*cf.* Smelror *et al.*, 2019) was described by Rawson (1982) from Kong Karls Land in the southern part of the Svalbard Archipelago. Such a stratigraphical interpretation of the Boreal ammonite faunas of *Arcticoceras* as corresponding to the Lower Bathonian (originally interpreted as indicative of the Boreal Middle Bathonian, *e.g.*, Callomon, 2003), results from recent studies of the ammonite succession in the Saratov area, central part of the Russian Platform (Mitta *et al.*, 2014, and earlier papers of the authors cited therein).

The stratigraphical interpretation of the middle part of the succession in the studied core shows that sedimentation has been very slow and condensed at least from the uppermost Callovian, through the whole Oxfordian, and a large part of the Kimmeridgian (possibly up to the lower part of the Elegans Zone). The total thickness of these deposits cannot be accurately determined, but it is between 12.65 and 16.57 m. The succession is possibly incomplete, with a probable stratigraphical gap including the upper part of the Middle Oxfordian, as well as a part of the Upper Oxfordian (see chapter on ammonite stratigraphy). It is also worth noting the occurrence of the stratigraphical interval between 68.56 and 63.28 m devoid of ammonites and with a poor bivalve fauna. This suggests the temporary existence of unfavorable environmental conditions for the development of the ammonite fauna. The total thickness of the deposits is markedly smaller than in Spitsbergen where coeval deposits may attain even about 70 metres in thickness (Koevoets *et al.*, 2018).

The upper part of the succession corresponding mostly to the Elegans Zone is more uniform and relatively thick. The deposits of the Elegans Zone (at least partly the equivalent of the Eudoxus Zone) correspond to the transgressive deposits of the younger part of the Hekkingen Formation formed on the western Barents Sea shelf during times of maximum flooding (see *e.g.*, Smelror *et al.*, 2001).

The described succession from core 7533/3-U-1 shows marked similarities in the general development of facies and the ammonite successions to those known from East Greenland and the shelf off western and mid Norway, as well as the North Sea area. The formation of the successions was governed by the rifting which appeared in Late Bajocian to Middle Bathonian and intensified in the Late Oxfordian and Kimmeridgian (Surlyk, 2003, and earlier papers cited therein). The rifting produced elevated fault-blocks, such as the one corresponding to the middle part of the succession studied in the core. The tectonic processes and the climatically-

controlled sedimentary cycles resulted also in the marine transgressions, peaking in the Late Kimmeridgian Eudoxus Chron (Elegans Chron). The recovered ammonites from the Upper Bajocian to the Upper Kimmeridgian comprise mostly Boreal Cardioceratidae, but the incoming of Subboreal Aulacostephanidae at two Kimmeridgian intervals offers possibilities to gain insight into fluctuations between the Boreal and Subboreal provinces during the tectonically active Late Jurassic time.

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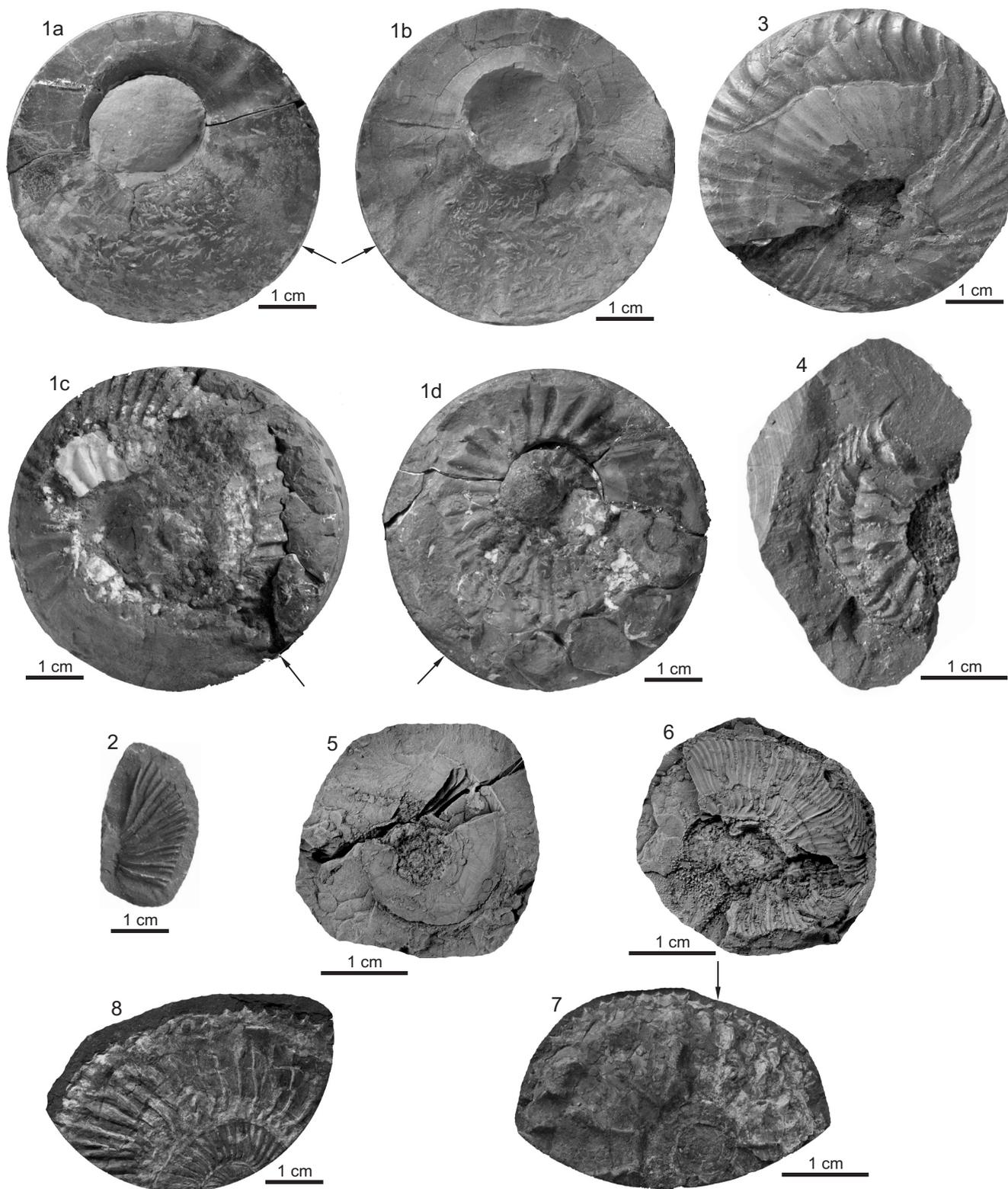
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Plates

PLATE 1

- Fig. 1a–d. *Cranocephalites indistinctus* Callomon, depth 117.24 m; a, b – final fragment of the phragmocone, and the body-chamber (with its outer part cut-off during drilling), inner mould (a) and outer mould (b); c, d – phragmocone, showing the ornamentation in its inner part, inner mould (c), and outer mould (d); Upper Bajocian, Indistinctus Zone; phragmocone/body-chamber boundary is arrowed; specimen no. MWG UW ZI/98/01
- Fig. 2. *Cranocephalites* sp., depth 115.67 m; fragment of inner whorls; Upper Bajocian; specimen no. MWG UW ZI/98/02
- Fig. 3. *Longaeviceras* ex gr. *L. placenta* (Leckenby) – *L. longaevum* (Leckenby), depth 75.22 m; Upper Callovian, Athleta Zone; specimen no. MWG UW ZI/98/05
- Fig. 4. *Cardioceras* (*Cardioceras*) *cordatum* (Sowerby), depth 70.99 m; Lower Oxfordian, Cordatum Zone – Cordatum Subzone; specimen no. MWG UW ZI/98/06; $\times 1.5$
- Fig. 5. *Cardioceras* (*Plasmatoceras*) *tenuistriatum* Borissjak, depth 70.37 m; Middle Oxfordian, Densiplicatum Zone; specimen no. MWG UW ZI/98/09 ; $\times 1.5$
- Fig. 6. *Cardioceras* (*Plasmatoceras*) *tenuicostatum* (Nikitin), depth 70.80 m; Middle Oxfordian, Densiplicatum Zone; specimen no. MWG UW ZI/98/08; $\times 1.5$
- Fig. 7. *Cardioceras* (*Subvertebriceras*) *zenaidae* Ilovaisky, depth 68.56 m; Middle Oxfordian, Densiplicatum Zone; phragmocone/body-chamber boundary is arrowed; specimen no. MWG UW ZI/98/10; $\times 1.5$
- Fig. 8. *Amoebites mesezhnikovi* (Sykes et Surlyk), depth 62.53 m; Lower Kimmeridgian, Kitchini Zone, Subkitchini Subzone; specimen no. MWG UW ZI/98/13

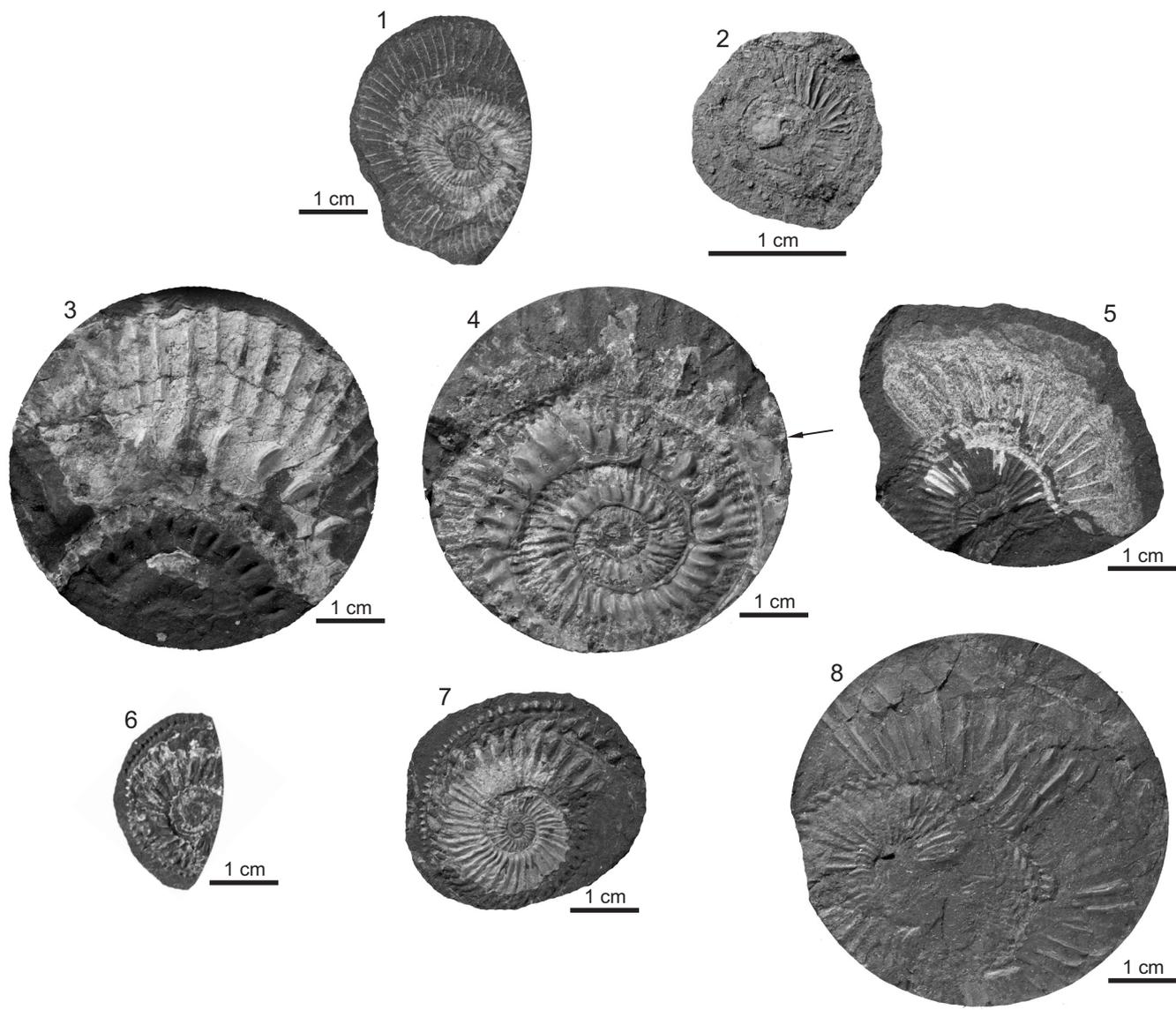


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PLATE 2

- Fig. 1. *Rasenia* sp., depth 61.41 m; Lower Kimmeridgian, Cymodoce Zone; specimen no. MWG UW ZI/98/15
- Fig. 2. *Euprionoceras* cf. *norvegicum* (Wierzbowski), depth 60.84 m; Kimmeridgian, Kochi Zone, *E. norvegicum* horizon; specimen no. MWG UW ZI/98/16; ×2
- Fig. 3. *Zenostephanus sachsii* (Mesezhnikov), depth 59.98 m; Upper Kimmeridgian, Mutabilis Zone, Mutabilis Subzone; specimen no. MWG UW ZI/98/17
- Fig. 4. *Zenostephanus scoticus* (Arkel et Callomon), depth 58.65 m; the phragmocone/body-chamber boundary is arrowed; Upper Kimmeridgian, Mutabilis Zone, Mutabilis Subzone; specimen no. MWG UW ZI/98/20
- Fig. 5. *Hoplocardioceras elegans* (Spath), depth 57.66 m; Upper Kimmeridgian, Elegans Zone; specimen no. MWG UW ZI/98/21
- Fig. 6. *Hoplocardioceras decipiens* (Spath), depth 56.83 m; Upper Kimmeridgian, Elegans Zone; specimen no. MWG UW ZI/98/24
- Fig. 7. *Hoplocardioceras elegans* (Spath) with heavy ribbing transitional to *H. decipiens* (Spath), depth 54.84 m; phragmocone; Upper Kimmeridgian, Elegans Zone; specimen no. MWG UW ZI/98/27
- Fig. 8. *Hoplocardioceras elegans* (Spath), depth 48.70 m; Upper Kimmeridgian, Elegans Zone; specimen no. MWG UW ZI/98/34

All specimens in natural size unless specified otherwise



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